

川东北元坝地区三叠系飞仙关组二段碳酸盐岩颗粒滩储层孔隙类型及其发育控制因素

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摘要:四川盆地东北部元坝地区三叠系飞仙关组二段(飞二段)发育多期碳酸盐岩颗粒滩储层,储层非均质性强,孔隙类型及发育控制因素不明,严重制约了储层评价与甜点预测。应用岩心、薄片、物性参数、成像测井以及压汞测试分析等资料,开展飞二段颗粒滩储层孔隙类型及主控因素研究,以期对飞二段有利储层预测提供科学依据。研究表明:①飞二段发育3期碳酸盐岩颗粒滩储层,包括鲕粒灰岩、豆粒灰岩以及砂屑灰岩等,部分储层发育白云石化作用形成的残余颗粒白云岩。②储层主要表现出低-中孔、低渗特征,可划分为裂缝型、孔隙型和裂缝-孔隙型3类。孔隙以颗粒内溶孔和铸模孔为主,粒间孔与晶间孔次之;裂缝则发育高角度构造裂缝和缝合线成岩缝。③孔隙发育主要受沉积条件、成岩作用和构造作用的共同控制。在不同沉积条件下发育的沉积微相和储层岩石类型决定了原生孔隙与胶结物的发育特征,并间接影响次生孔隙的发育程度;不同成岩阶段的环境差异控制了成岩作用的类型,进而制约了孔隙的类型与发育规模。其中,同生期-准同生期溶蚀作用、埋藏溶蚀作用和白云石化作用是飞二段孔隙发育的关键成岩作用类型;构造作用引起裂缝及其附近溶孔、溶洞的发育,大大提高了储层的孔、渗性能。④溶孔-裂缝型白云石化颗粒滩储层是飞二段最有利的储层类型,这得益于裂缝、溶孔和晶间孔的有效组合,使其孔、渗性能得到了显著改善。

关键词:孔隙类型;储层物性;碳酸盐岩颗粒滩储层;飞仙关组二段;元坝地区;川东北

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Pore types of granular shoal carbonate reservoirs in the 2nd member of the Triassic Feixianguan Formation, Yuanba area, northeastern Sichuan Basin and controlling factors for their development

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Abstract: The 2nd member of the Feixianguan Formation (also referred to as the Fei 2 Member, T_{1f}^{2}) in the Yuanba area, northeastern Sichuan Basin develops multi-stage granular shoal carbonate reservoirs characterized by strong heterogeneity. However, their pore types and controlling factors for development remain poorly understood, significantly restricting reservoir evaluation and sweet-spot prediction. In this study, by integrating core and thin-section observations, physical property parameters, formation micro-imager (FMI) logs, and high-pressure mercury injection (HPMI) analysis results, we investigate the pore types of granular shoal carbonate reservoirs in the Fei 2 Member and determine their major controlling factors for development, aiming to provide a scientific basis for predicting favorable reservoirs in this member. The research results indicate that three-stage granular shoal carbonate reservoirs occur in the Fei 2 Member, including oolitic limestones, pisolitic limestones, and calcarenites, as well as residual dolarenites derived from localized dolomitization. These reservoirs are characterized by low to moderate porosity and low

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permeability and can be classified into fractured, vuggy, and fractured-vuggy types. The pore system in the reservoirs is dominated by intragranular dissolution pores and moldic pores, followed by intergranular and intercrystalline pores. Fractures are mainly of high-angle and stylolite-related types. The pore development is jointly controlled by sedimentary conditions, diagenesis, and tectonism. Both sedimentary microfacies and reservoir lithologies resulting from different sedimentary settings determine the characteristics of primary pores and cements, while indirectly governing the extent of secondary pore development. Environmental variations at diverse diagenetic stages control the types of diagenesis, thereby constraining the pore development in terms of type and scale. Among various diagenesis types, syngenetic-penecontemporaneous dissolution, burial dissolution, and dolomitization are critical to pore development in the member. Tectonic activity resulted in the formation of fractures, as well as adjacent dissolution pores and vugs, thus significantly enhancing reservoir porosity and permeability. Dolomitized granular shoal reservoirs of the dissolution pore-fracture type are identified as the most favorable reservoir type in the member, as their porosity and permeability are substantially enhanced by the effective combination of fractures, dissolution pores, and intercrystalline pores.

Key words: pore type, reservoir physical property, granular shoal carbonate reservoir, the 2nd member of Feixianguan Formation, Yuanba area, northeastern Sichuan Basin

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四川盆地东北部环开江-梁平海槽台缘带下三叠统飞仙关组(T_{1f})发育了多套碳酸盐岩颗粒滩相储层。自20世纪90年代该层系获得工业气流以来,已相继发现渡口河、罗家寨、滚子坪及普光等大-中型气藏。近年来,海槽西侧的九龙山地区亦取得勘探突破,进一步揭示了下三叠统飞仙关组颗粒滩相气藏巨大的资源潜力^[1-14]。与之相邻的元坝地区勘探同样活跃,2007年,河坝1井在飞仙关组三段(飞三段, T_{1f}^3)和下三叠统嘉陵江组二段(嘉二段, T_{1j}^2)获得重大突破并投入试采;此后部署的元坝1井等多口探井,均在长兴组与飞(飞仙关组)二段见到良好油气显示^[15-18]。其中,长兴组礁滩相储层一直是关注的焦点,相关学者普遍认为其储集性能远优于飞仙关组储层^[19-21]。前人对开江-梁平海槽东侧的飞仙关组颗粒滩相储层进行了大量研究,涉及层序地层、沉积相、成岩作用、储层特征和成藏条件等诸多方面^[5,9,22-33],而针对海槽西侧飞二段的颗粒滩相储层研究则相对比较匮乏。近年来,元坝27井和元坝204井飞仙关组颗粒滩相储层获得了工业气流,这一发现推动了元坝地区飞仙关组颗粒滩相储层研究的深入^[19,34]。元坝地区飞二段发育多期颗粒滩,而颗粒滩相储层的沉积成因复杂,储层厚度横向变化大,且不同类型的颗粒滩相储层在岩石学特征、物性特征和孔隙类型方面存在显著差异,表现出极强的非均质性,且不同类型孔隙和有利储层发育的控制因素不明,有利

储层预测难度大,制约了元坝地区飞仙关组的有效勘探和开发^[35]。本文利用岩心、薄片、物性参数、压汞测试和成像测井等资料,对元坝地区飞二段颗粒滩相储层开展岩石学特征、物性特征和孔隙类型及结构特征的研究,探讨不同类型孔隙和有利储层发育的主控因素,以期为元坝地区飞仙关组下一步的勘探和开发部署提供科学依据。

1 研究区概况

元坝地区位于四川盆地东北部的开江-梁平海槽西侧(图1),构造上处于川东高陡构造带,处于川北坳陷与川中平缓构造带的交汇地带^[36-37]。这一特殊的构造位置使该区成为盆地内应力与构造形变的调节区,其西北、东北和南部分别与九龙山构造带、通南巴构造带和川中低缓构造带相接。元坝地区飞仙关组的碳酸盐岩颗粒滩相沉积明显受开江-梁平海槽演化的控制。开江-梁平海槽形成于晚二叠世—早三叠世初期的被动陆缘拉张背景,表现为一个北西宽、南东窄的狭长三角状断陷海槽。元坝地区二叠纪—三叠纪的沉积环境受海槽的形成与关闭、米仓山-大巴山隆起以及峨眉山玄武岩喷发等重大构造-岩浆事件的协同影响^[40]。东吴运动使得开江-梁平演变为陆棚环境,在元坝、龙岗和普光等地区广泛发育长兴组生物礁,近年

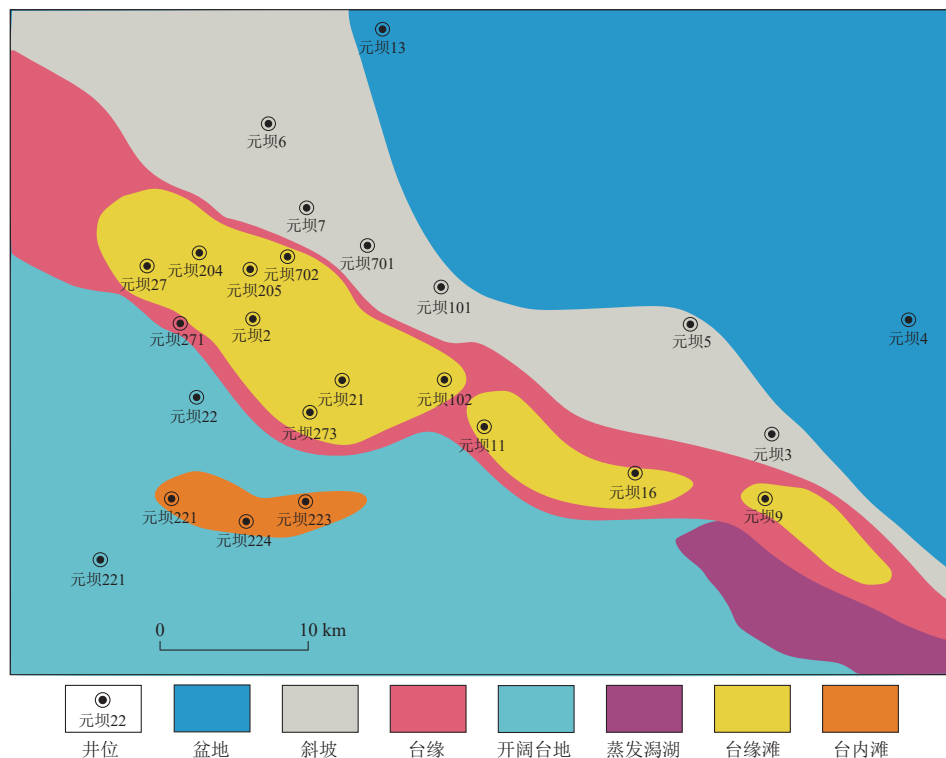


图3 川东北元坝地区飞二段沉积中期沉积相平面分布

Fig. 3 Map showing the distribution of mid-term sedimentary facies in the T_1f^2 , Yuanba area, northeastern Sichuan Basin

3 颗粒滩储层物性特征

对飞二段不同样品的物性统计(图7)及对应的直方图(图8)分析显示,孔隙度主要分布于2.01%~16.27%,平均值为4.25%。孔隙度频率呈双峰式分布,孔隙度主峰位于2.00%~4.00%(占样品的58.0%),次峰为4.00%~6.00%(占样品的29.3%),两者合计占比超过87.0%,孔隙度大于6.00%的样品仅占12.7%,整体属于低-中孔储层。渗透率介于 $(0.002 \sim 1\,340.284) \times 10^{-3} \mu\text{m}^2$,平均值仅为 $0.110 \times 10^{-3} \mu\text{m}^2$,分布极不均匀。约70.0%的样品渗透率低于 $0.100 \times 10^{-3} \mu\text{m}^2$,其中以渗透率为 $(0.010 \sim 0.100) \times 10^{-3} \mu\text{m}^2$ 的样品占比最高(42.5%),而渗透率高于 $10.000 \times 10^{-3} \mu\text{m}^2$ 的样品仅占3.3%(图8)。

进一步对比3期鲕粒滩的物性可见显著阶段性差异:第二期滩体物性最佳,平均孔隙度与平均渗透率分别为4.05%和 $0.186 \times 10^{-3} \mu\text{m}^2$;第一期与第三期物性相对较差,平均孔隙度分别为3.08%和3.18%,平均渗透率分别为 $0.098 \times 10^{-3} \mu\text{m}^2$ 和 $0.137 \times 10^{-3} \mu\text{m}^2$ 。

储集空间和渗流通道的配置关系是储层类型划分的根本依据。在碳酸盐岩储层中,裂缝空间分布复杂且不均,其与基质孔隙的有机结合,形成了特有的双重

孔隙结构系统^[63-65]。本研究基于孔-渗关系(图9)与铸体薄片观察结果,将飞二段储层划分为孔隙型、裂缝型及裂缝-孔隙型3类,其孔-渗耦合关系差异显著。其中,裂缝-孔隙型储层是区内最重要的储集类型,以次生孔隙为主要储集空间(平均孔隙度为4.31%),微裂缝则显著提升了孔隙连通性,形成了孔-渗协同增长的典型特征,如元坝204井因裂缝发育,其渗透率最高可达 $1\,340.284 \times 10^{-3} \mu\text{m}^2$;孔隙型储层兼具原生与次生孔隙,但因缺乏有效裂缝系统,渗流能力有限,多表现为高孔、低渗特征;裂缝型储层则以裂缝网络主导储层渗流,呈现低孔、中-高渗特征,孔隙度与渗透率之间的相关性微弱。

4 颗粒滩储层孔隙类型与结构特征

4.1 孔隙类型及其特征

元坝地区飞二段颗粒滩储层的储集空间类型主要为孔隙和裂缝。孔隙以粒内溶孔与铸模孔为主,粒间孔较少见;在残余砂屑白云岩中发育晶间孔(图3)。裂缝以构造裂缝和成岩缝为主(图3a,图9,图10)。

粒内溶孔多发育于飞二段鲕粒灰岩中(图3),在砂屑灰岩和豆粒灰岩中则较为少见。在鲕粒灰岩发育的

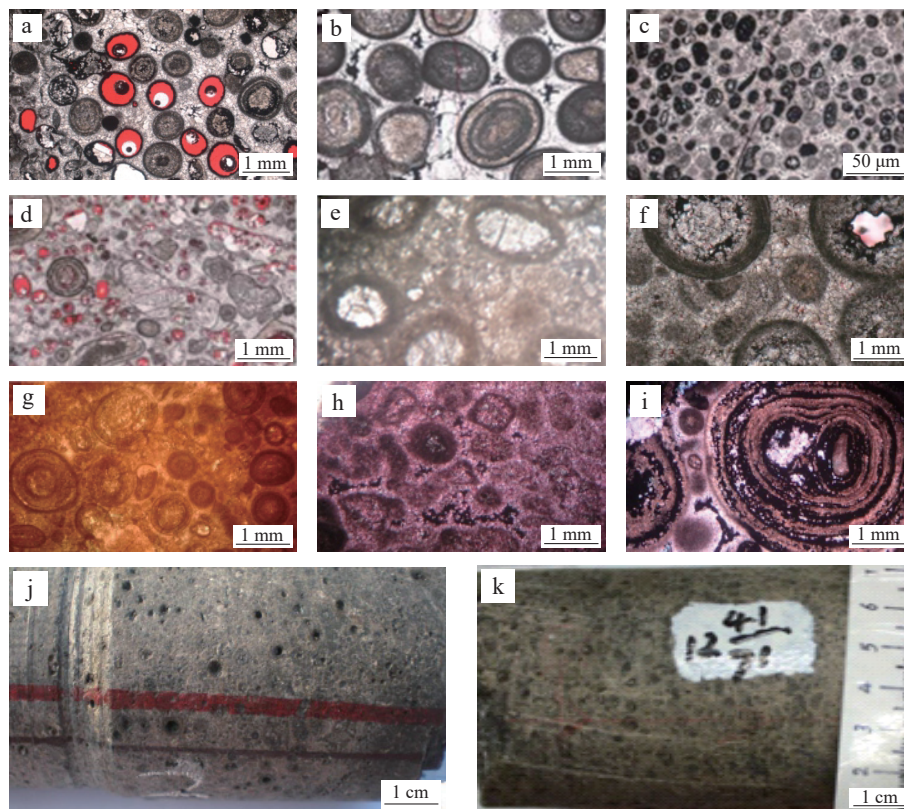


图4 川东北元坝地区飞二段鲕粒灰岩和豆粒灰岩特征镜下照片

Fig. 4 Images showing the characteristics of oolitic and pisolitic limestones in the T_1f^2 , Yuanba area, northeastern Sichuan Basina.

a. YB27井,埋深6 125.81 m,亮晶鲕粒灰岩,铸体薄片,单偏光(-); b. YB21井,埋深6 660.28 m,亮晶鲕粒灰岩,铸体薄片,单偏光(-); c. YB27井,埋深6 126.67 m,鲕粒灰岩,铸体薄片,单偏光(-); d. YB27井,埋深6 131.97 m,溶孔鲕粒灰岩,铸体薄片,单偏光(-); e. YB9井,埋深6 645.00 m,泥晶鲕粒灰岩,铸体薄片,单偏光(-); f. YB22井,埋深6 408.49 m,亮晶鲕粒灰岩,孔隙内壁残留少量有机质,铸体薄片,单偏光(-); g. YB22井,埋深6 444.00 m,亮晶鲕粒灰岩,普通薄片(染色),单偏光(-); h. YB101井,埋深6 795.00 m,鲕粒砂屑灰岩,普通薄片(染色),单偏光(-); i. YB204井,埋深6 416.05 m,含豆粒鲕粒灰岩,普通薄片(染色),单偏光(-); j. YB27井,埋深6 127.92 ~ 61 28.10 m,溶孔鲕粒灰岩,岩心照片; k. YB204井,埋深6 422.11 ~ 6 422.24 m,溶孔鲕粒灰岩,岩心照片

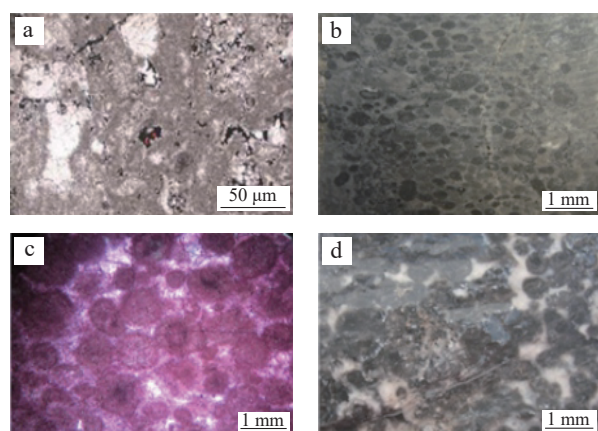


图5 川东北元坝地区飞二段砂屑灰岩特征镜下照片

Fig. 5 Thin-section photomicrographs of calcarenites from the T_1f^2 , Yuanba area, northeastern Sichuan Basin

a. YB271井,埋深6 208.12 m,藻砂屑灰岩,普通薄片,单偏光(-); b. YB101井,埋深6 797.70 m,亮晶砂屑灰岩,普通薄片,单偏光(-); c. YB101井,埋深6 793.80 m,亮晶鲕粒砂屑灰岩,普通薄片(染色),单偏光(-); d. YB101井,埋深6 797.70 m,亮晶含鲕砂屑灰岩,普通薄片,单偏光(-)

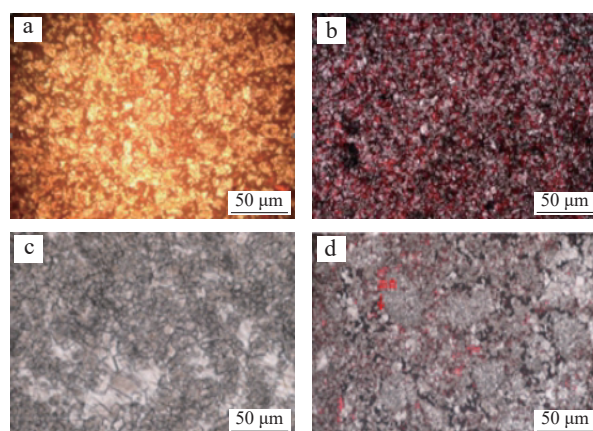


图6 川东北元坝地区飞二段残余砂屑白云岩特征镜下照片

Fig. 6 Thin-section photomicrographs of residual doloarenites from the T_1f^2 , Yuanba area, northeastern Sichuan Basin

a. YB21井,埋深6 656.20 m,残余砂屑灰质白云岩,普通薄片(染色),单偏光(-); b. YB21井,埋深6 657.09 m,残余砂屑灰质白云岩,普通薄片,正交光(-); c. YB9井,埋深6 681.77 m,残余砂屑含灰质白云岩,普通薄片,单偏光(-); d. YB9井,埋深6 679.26 m,溶孔残余砂屑粉晶白云岩,铸体薄片,单偏光(-)

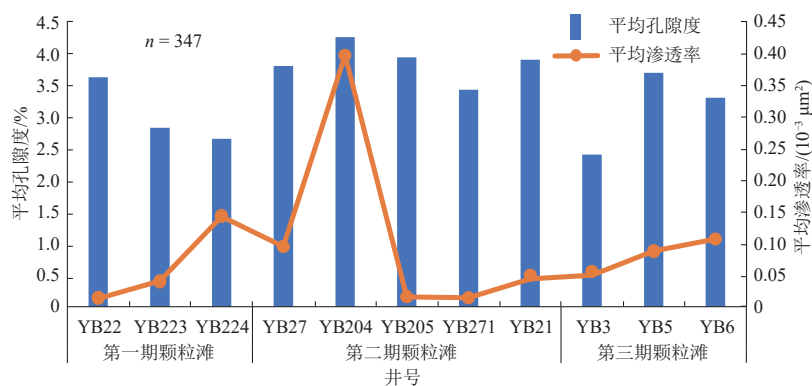


图7 川东北元坝地区飞二段储层物性统计柱状图

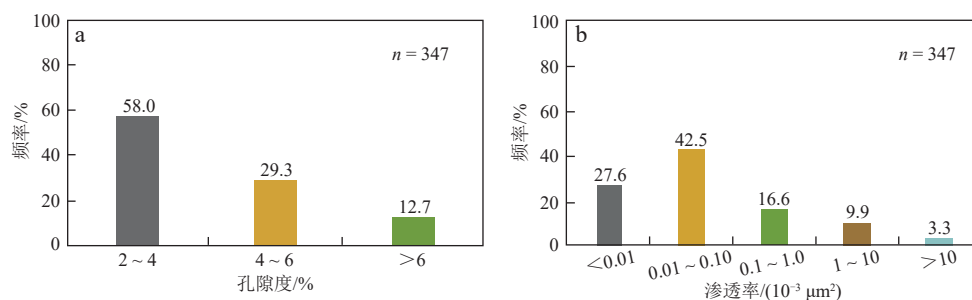
Fig. 7 Bar chart showing the reservoir physical properties of the T_1f^2 , Yuanba area, northeastern Sichuan Basin

图8 川东北元坝地区飞二段储层孔隙度(a)与渗透率(b)频率分布直方图

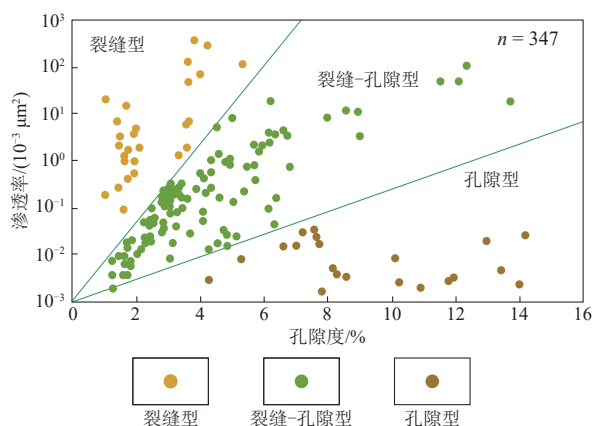
Fig. 8 Frequency distribution histograms of porosity (a) and permeability (b) in the T_1f^2 , Yuanba area, northeastern Sichuan Basin

图9 川东北元坝地区飞二段储层孔隙度与渗透率的关系

Fig. 9 Porosity-permeability cross plot of the T_1f^2 , Yuanba area, northeastern Sichuan Basin

岩心表面,常可见密集分布的粒内溶孔,多由鲕粒的选择性溶蚀作用形成。根据成因,研究区鲕粒灰岩的储集空间可分为3类:①选择性溶孔,主要为鲕粒内溶孔,由早期不稳定矿物(如文石)溶解形成;若鲕粒完全溶蚀仅保留外壳,则形成鲕模孔(图3c,m),此类孔隙早期多被沥青充填,晚期则多被矿物部分充填,呈半月状。②埋藏溶孔,以粒间溶孔为主(图3f),由后期深埋环境下胶结物发生溶蚀形成,属于非选择性溶蚀。③残余原生

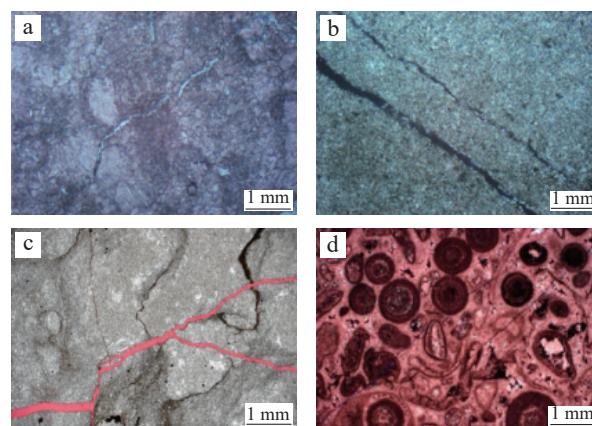


图10 川东北元坝地区飞二段裂缝类型及其特征镜下照片

Fig. 10 Images showing the fracture types and characteristics in the T_1f^2 , Yuanba area, northeastern Sichuan Basin

- a. YB4井,埋深6 673.00 m,云质残余砂屑灰岩,铸体薄片,单偏光(-);
 b. YB4井,埋深7 148.00~7 150.00 mm,泥晶灰岩,普通薄片,单偏光(-);
 c. YB3井,埋深6 703.38 m,微晶灰岩,铸体薄片,单偏光(-);
 d. YB204井,埋深6 411.08 m,亮晶鲕粒灰岩,普通薄片(染色),单偏光(-)

孔,即粒间孔,本身不发育且形态不规则。另外,沥青的分布是关键成岩时间标志:充填于早期鲕模孔中的沥青表明该部分孔隙形成于油气充注之前,属于有效孔隙。晶间孔主要发育于残余砂屑白云岩中。白云石化

作用后,白云石以粉晶结构为主,晶间孔均匀发育。部分白云石晶体发生溶蚀可形成晶粒溶孔(图6d),局部被沥青充填或半充填,面孔率一般介于3%~10%。

飞二段颗粒滩中发育的裂缝主要包括构造裂缝和成岩缝(图3a,图4a,图10)。裂缝系统对储层渗流具有关键控制作用。元坝地区飞仙关组发育3期构造裂缝^[23-26,42,66],主要分布在靠近九龙山背斜一侧,以高角度裂缝为主,可分为2期:早期裂缝多被方解石、白云石和石英充填而呈闭合状;晚期裂缝缝壁洁净、未见矿物充填,缝宽介于0.02~0.50 mm(图10),是当前有效的渗流通道。此外,由压溶作用形成的缝合线在颗粒滩中主要沿颗粒边缘分布,且在缝合线上常见沥青浸染(图11),指示其曾是埋藏期溶蚀流体与烃类运移的优势通道,构成一类易被忽视却至关重要的储集空间。此套“构造-成岩”双成因裂缝系统控制了储层的非均质性,并形成了复杂的渗流通道。基于薄片鉴定、扫描电镜及成像测井资料分析,可将飞二段裂缝分为高角度高导缝(倾角>45°)、低角度高导缝(倾角≤45°)、微裂缝、诱导缝和缝合线,未见低导缝。其中,高角度高导缝数量多于低角度高导缝;微裂缝在元坝9井中较为发育,诱导缝则在元坝223井中较为发育。

4.2 孔隙结构特征

元坝地区飞二段储层的储集性能主要受孔隙-喉道系统制约。压汞资料显示,储层孔隙结构具有较强的非均质性,主要参数如下:①排驱压力(p_{c10})介于0.086 2~44.871 6 MPa,算术平均值为6.998 9 MPa,

几何平均值为3.370 8 MPa,几何平均值显著低于算术平均值,表明储层喉道连通性呈正偏态分布;②中值压力(p_{c50})变化范围为0.192 7~180.085 8 MPa,算术平均值为29.548 0 MPa,几何平均值为19.384 0 MPa,较大的数值跨度进一步印证了孔隙结构的非均质性;③孔隙中值喉道半径(R_{c50})最小值仅为0.004 2 μm ,最大值可达3.891 2 μm ,算术平均值为0.109 6 μm ,几何平均值为0.040 6 μm ,几何平均值远小于算术平均值,表明储层喉道以微喉道为主,且分布离散性强;④中值喉道半径大于0.075 0 μm 的样品占比介于0.131 2%~87.995 7%,算术平均值为30.504 1%,几何平均值为17.898 8%,较大的波动范围表明不同样品有效连通喉道的发育程度差异显著。总体而言,飞二段储层的储集空间以小孔和微孔(孔径<2 μm)为主,喉道则以细喉为主,粗-中喉样品合计仅占3.480 0%,孔-喉组合多呈“中孔-细喉”特征。与之相应,较高的排驱压力与中值压力直接反映出孔隙结构整体较差,流体渗流所需克服的阻力较大。

5 颗粒滩储层孔隙发育控制因素

元坝地区飞二段不同颗粒滩储层在孔、渗性能上存在显著差异,主要受沉积条件、成岩作用和构造作用的共同控制^[52,59-64,67-68]。结合研究区实际资料分析,这对储层发育的控制作用可概括如下:①沉积条件决定了颗粒滩储层的岩石类型及特征,直接控制原生孔隙和胶结物的发育,进而间接影响次生孔隙的发育程度;②不同成岩阶段的成岩环境差异控制成岩作用的类型与强度,间接决定孔隙的类型及演化;③构造作用则控制裂缝的发育程度,并进一步促进附近溶孔-溶洞的形成,改善储层渗流能力。

5.1 沉积条件控制储层岩石类型及孔隙发育

沉积相带和储层岩石类型是储层发育的基础。沉积条件不仅决定了颗粒滩的岩石组分与结构,直接制约原生孔隙与胶结物的发育,还通过影响早期成岩环境间接控制次生孔隙的发育程度。

元坝地区飞二段颗粒滩储层主要形成于台内滩和台缘滩两种沉积环境^[9,23,30,69]。不同环境中形成的颗粒滩储层在岩石类型、物性特征及孔隙类型方面均存在明显不同,表现出显著的相控分异特征。根据岩心物性数据统计,飞二段储层孔隙度主体分布在2.00%~26.80%。其中,台缘滩储层平均孔隙度为3.84%,明显优于台内滩储层的2.95%,表明台缘带储层物性显

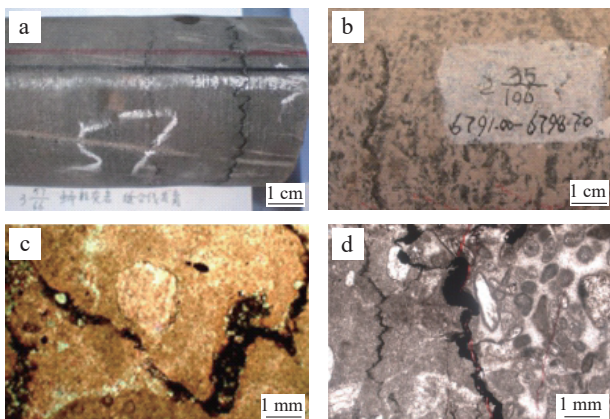


图11 川东北元坝地区飞二段缝合线类型及其特征照片

Fig. 11 Images showing the stylolite types and characteristics in the T_1f^2 , Yuanba area, northeastern Sichuan Basin

a. YB27井,埋深6 126.51~6 126.57 m,泥晶灰岩,岩心照片;b. 砂屑灰岩,YB101井,埋深6 794.80~6 794.87 m,岩心照片;c. YB27井,埋深6 115.50 m,泥晶灰岩,普通薄片镜下照片,正交光(-);d. YB27井,埋深6 133.80 m,砂屑灰岩,高角度缝合线,铸体薄片镜下照片,单偏光(-)

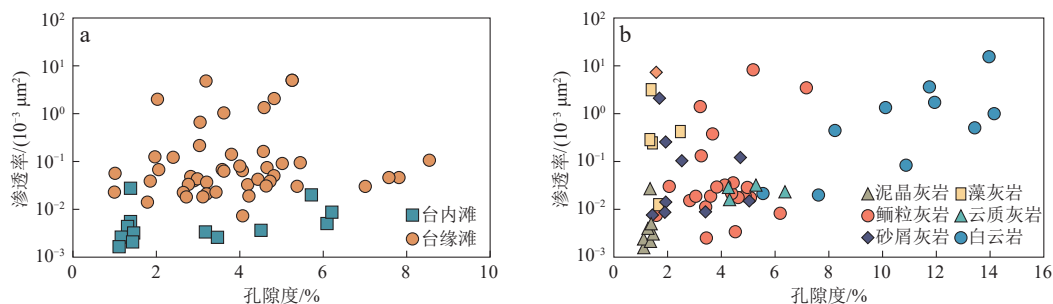


图12 川东北元坝地区飞二段不同相带(a)及不同岩性(b)储层孔-渗相关性

Fig. 12 Correlation between porosity and permeability for different facies zones (a) and different lithologies (b) in the T_1f^2 reservoirs, Yuanba area, northeastern Sichuan Basin

著优于台内滩。此外,同一滩体内部不同部位的储集条件亦存在明显差异,储层质量自滩核向滩缘过渡带及滩缘依次变差。岩性对物性的控制作用同样显著。根据不同岩性的孔-渗散点图(图12)可知,颗粒白云岩的物性最好,其次为鲕粒灰岩、砂屑灰岩和豆粒灰岩。颗粒白云岩之所以表现出最优的储集性能,应与其经历了白云石化作用和埋藏溶蚀作用的双重改造密切相关。

受海平面变化旋回控制,元坝地区飞仙关组台地边缘与台内鲕粒滩主要发育于各级层序旋回的顶部。该时期水体整体变浅、水动力能量较强,颗粒滩储层粒度较粗,颗粒间悬浮的泥晶被淘洗干净,原生孔隙较发育,为后期次生孔隙的形成奠定了基础。台缘滩与台内滩多发育于古地貌高地^[70],这为同生-准同生期大气淡水的淋滤与溶蚀作用提供了有利条件,是早期孔隙发育的关键控制因素。此外,台缘滩邻近海槽生油区,有机质成熟过程中排放的有机酸先进入临近的台缘滩储层,溶解颗粒及早期亮晶胶结物,形成大量溶蚀次生孔隙。因此,台缘滩的储集性能普遍优于台内滩。碳酸盐岩储层的演化普遍遵循“马太效应”,即早期具备良好原生孔隙的储层,在后续成岩改造中往往占据优势,更易于形成规模大、连通性好且分布稳定的优质储集空间^[71-72]。颗粒滩单个滩体的厚度和分布面积同样控制着原生孔的发育与保存,进而影响早期胶结物的发育程度及压实作用强度。单个滩体规模越大,次生粒间孔的发育与保存机率也就越高。

5.2 成岩环境差异控制孔隙类型及发育程度

元坝地区飞二段颗粒滩储层经历了同生-准同生、早成岩、中成岩和晚成岩等多个阶段的成岩改造。不同埋藏阶段的成岩环境差异显著,导致成岩作用类型各异,进而控制了孔隙的类型与发育程度(图13)。其中,同生-准同生期溶蚀作用、埋藏溶蚀作用及白云

石化作用有利于孔隙发育,而压实作用、胶结作用和充填作用则显著降低了孔隙度。

元坝地区飞二段颗粒滩中发育的溶蚀孔隙以粒内溶孔和铸模孔为主(图3),具有明显的选择性溶蚀特征。部分孔隙呈新月状(或半月状),为同生期渗流带大气淡水淋滤与选择性溶蚀的典型标志^[73];与之相对应,这些孔隙充填物呈现昏暗-不发光的阴极发光特征,指示低离子浓度、低氧化性的大气水成岩环境^[42,74-77]。上述特征共同限定了该期建设性成岩作用的发生时期与流体性质。部分粒内溶孔在垂向上被不同结构的胶结物或泥粉晶充填,并发育典型的示顶底构造,表明其形成后经历了渗流成岩环境。发育粒内溶孔的颗粒滩储层,压实弱、颗粒多呈漂浮或点接触,集中分布于高频层序的中、上部。这些特征共同指示其形成时间较早,与高频海平面下降引起的同生期大气淡水淋滤溶蚀作用密切相关,属于典型的层序控储型孔隙^[73]。与之相对,少量粒间溶孔可能源于埋藏阶段酸性流体的二次溶蚀,该期流体沿早期孔隙网络运移,对储层进行叠加改造。综上所述,研究区孔隙系统呈现“同生期层序控制型溶蚀奠定基础,埋藏期流体改造型溶蚀叠加优化”的阶段性建设模式。

白云石化作用是控制飞二段优质储层发育的关键成岩作用类型。研究区白云岩与灰岩的储集性能差异显著(图12),白云岩物性优势明显,约60%的样品孔隙度超过2%,而灰岩储集物性普遍较差。单井数据显示,储层厚度与鲕粒白云岩厚度呈显著正相关,但与鲕粒灰岩厚度无明显关联,表明白云岩构成了飞二段有利储层的物质基础^[3,78]。综合来看,白云石化与非白云石化颗粒滩储层的物性差异,主要取决于埋藏酸性流体侵入后的运移模式。在白云石化颗粒滩中,晶间孔和晶粒溶孔发育,流体易于发生侧向输运,进而形成更大规模的成岩改造储层(图13);而残余粒间孔发育

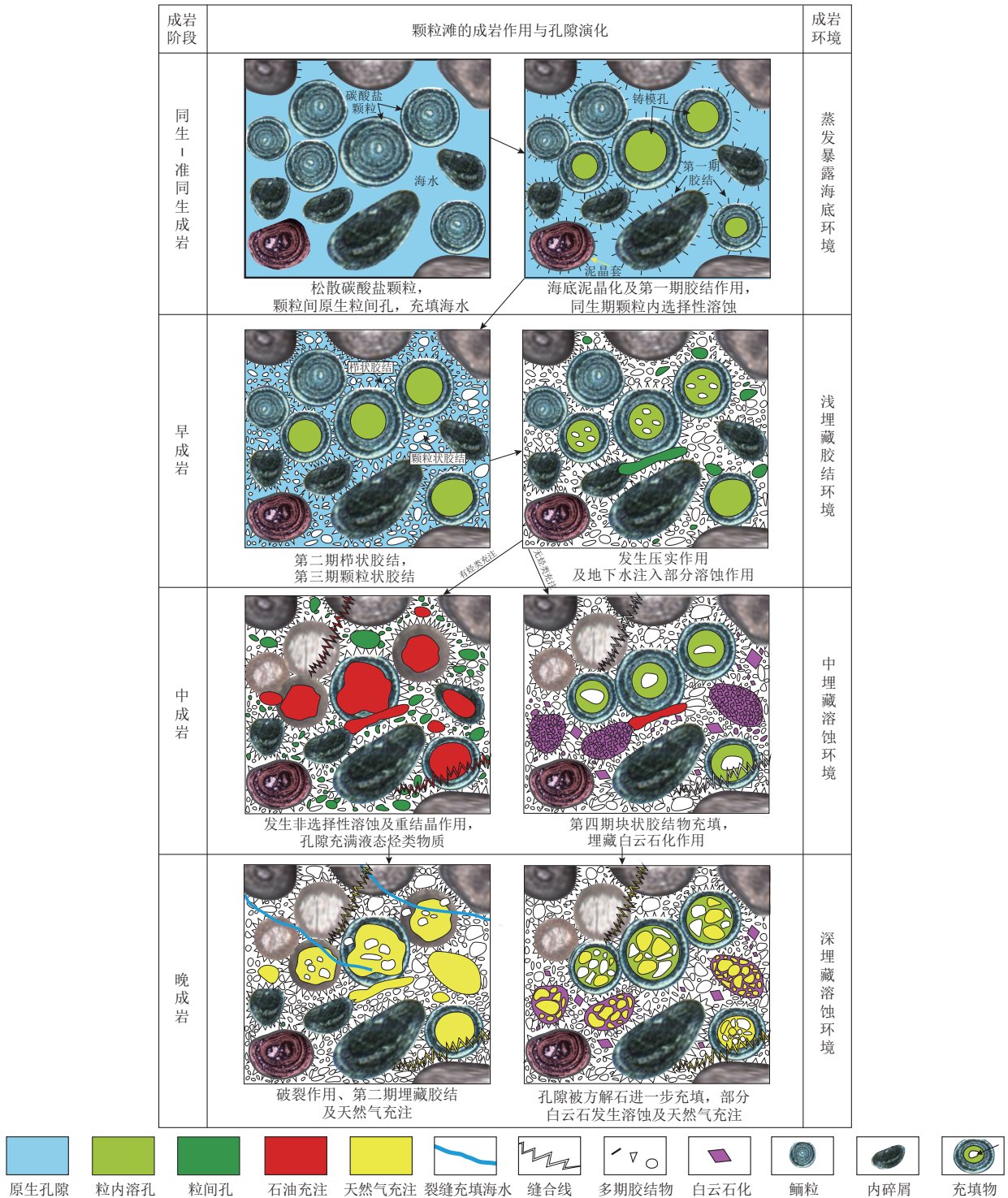


图 13 川东北元坝地区飞二段颗粒滩储层成岩及演化模式示意图

Fig. 13 Diagram showing the diagenesis and evolution model for granular shoal reservoirs in the T_1f^2 , Yuanba area, northeastern Sichuan Basin

的鲕粒灰岩储层,因孔隙孤立、缺乏有效侧向运移通道,流体活动主要局限于断裂-裂缝系统附近,其对储层的影响相对有限。此外,深埋藏条件下白云石的溶解速率明显高于方解石,这也是白云石化颗粒滩孔隙更发育的另一重要原因。

5.3 构造作用促进裂缝及溶蚀孔隙发育

构造作用对储层孔隙发育的影响主要体现在裂缝的形成与演化方面。受燕山期和喜马拉雅期陆内构造及新生代晚期青藏高原东缘冲断隆升的共同影响,研究

区发育了龙门山北段冲断带、米仓山南缘冲断带以及九龙山背斜。元坝气田西北部(如YB27井和YB204井)飞仙关组储层物性显著优于东部,部分原因在于其处于构造高部位,有位置优势^[79-80]。研究区裂缝系统受九龙山背斜控制,其多期活动清晰记录了构造—成岩—烃类充注的耦合过程^[81-85]。如前所述,飞二段共识别出3期裂缝:①第一期形成于早成岩阶段,充填沥青污染的方解石,被后期裂缝切穿;②第二期形成于中成岩A期,即液态烃阶段,部分充填沥青,切割第一期但被第三期切割;③第三期形成于晚成岩阶段(气态烃充注),基本无充填,切割前两期。其中,第二期与第三期裂缝对储层具有建设性贡献,尤以未充填的第三期裂缝贡献最大^[25]。这些裂缝不仅是重要的渗流通道,更是孔隙水与有机酸运移的优势路径,有效促进了裂缝周边的溶蚀扩孔作用,从而在改善储层渗透性的同时,局部提升了孔隙度。

6 结论

1) 元坝地区飞二段发育3期颗粒滩,包括台内滩和台缘滩,颗粒滩储层主要为鲕粒灰岩和砂屑灰岩,部分砂屑灰岩发生白云石化而成残余砂屑白云岩。飞二段储层主要以低孔—中孔、低渗为主,第二期鲕粒滩的物性及孔隙结构特征最优。按孔、渗空间配置关系,可将颗粒滩储层划分为孔隙型、裂缝型和裂缝—孔隙型3类,其中裂缝—孔隙型储层最为有利。

2) 飞二段颗粒滩储层储集空间类型包括孔隙和裂缝2类。孔隙以鲕粒内溶孔和铸模孔为主,少见粒间孔,在残余砂屑白云岩中发育晶间孔和晶粒溶孔。裂缝发育3期,以构造裂缝和成岩缝为主,其中构造裂缝主要分布在靠九龙山背斜一侧,以高角度裂缝为主;成岩缝以缝合线为主,多围绕颗粒边缘分布。研究区孔隙结构以中孔—细喉和小孔—细喉为主,整体孔隙结构与孔、渗条件较差。

3) 受沉积条件、成岩作用与构造作用的共同控制,飞二段不同颗粒滩储层的孔、渗特性差异显著。沉积条件所决定的沉积相带与岩石类型,控制了原生孔隙和胶结物的发育,进而间接影响次生孔隙的形成,其中颗粒云岩物性最优。同生—准同生期淋滤与溶蚀作用、埋藏溶蚀作用以及白云石化作用是飞二段颗粒滩储层孔隙发育的关键成岩过程。构造作用则通过裂缝发育及其附近溶孔、溶洞的形成,显著改善了储层的渗透性能。综合分析表明,溶孔—裂缝型白云石化颗粒滩储层是研究区最有利的储层类型。

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