

准噶尔盆地博格达地区中二叠统芦草沟组热液硅质结核特征及页岩油意义

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摘要:硅质岩已被广泛应用于重建古环境、示踪热液循环和硅化过程等方面,并且对陆相页岩油勘探开发具有重要意义。以准噶尔盆地博格达地区芦草沟组为研究对象,综合运用X射线衍射、元素面扫、主微量元素、镜质体反射率、岩石热解和总有机碳分析方法对芦草沟组硅质岩成因机制及油气意义进行了研究。结果表明:芦草沟组硅质结核中软沉积物变形构造非常发育,主要由隐-微晶石英组成,指示结核为同沉积期形成。结核内部发育重晶石、锰菱铁矿、镍铁矿和蓝铜矿等热液矿物,并且稀土元素具有总量低、重稀土元素相对富集和正Eu异常等特征,指示其为热液成因,说明二叠纪研究区存在间歇性热液活动。热液带来的营养元素和金属阳离子提高了古生产力和古盐度,促使底水还原性增强,有利于有机质富集。硅质岩中孔隙和裂缝发育,其中显著富集的石英提高了储层脆性,有利于压裂改造。硅质岩伴生的热液活动参与了有机质生烃演化,使页岩油黏度降低,可动性增强,有利于页岩油开发。

关键词:硅质结核;热液;有机质;古沉积环境;页岩油;芦草沟组;准噶尔盆地

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Characterizing hydrothermal siliceous nodules to guide shale oil exploration in the Middle Permian Lucaogou Formation, Bogda area, Junggar Basin

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Abstract: Siliceous rocks are frequently included in reconstructing paleoenvironment, tracing hydrothermal circulation and understanding silicification process, aiming at using the results to guide the exploration and development of continental shale oil. Siliceous rock samples from the Lucaogou Formation in Bogda area, Junggar Basin, were observed through X-ray diffraction and scanning electron microscope and analyzed with major and trace elements, vitrinite reflectance, Rock-Eval and total organic carbon data to identify their genetic mechanisms and hydrocarbon significance. The results show that the siliceous nodules in the samples host highly-deformed soft-sediments mainly composed of crypto-microcrystalline silica, indicating a synsedimentary product. Hydrothermal minerals such as barite, mangano-siderite, trevorite and azurite are developed inside the nodules with low total content of rare earth elements and relative high content of heavy rare earth elements as well as positive Eu anomalies, all indicating a hydrothermal origin and intermittent hydrothermal activities in the Bogda area during the Permian. The nutrients and metal cations, brought by the hydrothermal fluid, improved the paleo-productivity and paleo-salinity, and enhanced the reducibility of the bottom water, which was conducive to the enrichment of organic matter. Well-developed pores and fractures with enriched quartz greatly improved the

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brittleness of the rocks, making the reservoirs in the formation more responsive to fracturing. The hydrothermal activities associated with siliceous rocks participated in the evolution of hydrocarbon generation of organic matters, resulting in reduced crude viscosity and enhanced crude mobility which are conducive to shale oil development.

Key words: siliceous nodule, hydrothermal fluid, organic matter, ancient sedimentary setting, shale oil, Lucaogou Formation, Junggar Basin

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准噶尔盆地油气资源丰富^[1],其中芦草沟组是盆地内重要的烃源岩层系^[2],同时也是陆相页岩油勘探开发的重点层系^[3]。芦草沟组具有混合沉积的特征^[4],岩性复杂,陆源碎屑、火山–热液沉积物和碳酸盐矿物均有发育^[5]。其中,硅质岩作为一类特殊的化学沉积岩在准噶尔盆地和邻区三塘湖盆地芦草沟组中广泛分布^[6–7]。然而,硅质岩具有成因复杂且硅质来源多样等特征,在沉积期蒸发沉淀或生物诱导^[8]、成岩期交代作用下均可形成,硅质来源主要包括陆源碎屑、生物硅和深源流体等^[9]。目前,对于准噶尔盆地和三塘湖盆地二叠系硅质岩成因主要有以下几种认识:热液成因^[10–11]、火山灰成因^[12]、蒸发和生物诱导混合成因^[8]、热液和碱湖蒸发混合成因^[13]等。硅质来源和硅质岩沉积过程的复杂性增加了储层非均质性^[14],进一步加大了页岩油勘探开发难度^[15]。

硅质岩由于性质稳定、结构致密以及抗风化能力强等特点^[9],已经被广泛应用于重建古沉积环境^[16]、示踪古海洋热液循环及硅化过程^[17]。二叠纪,全球很多国家或地区如北美、菲律宾、意大利和中国均发生了硅质沉积的显著富集,被称为二叠纪硅质沉积事件。因此,二叠系硅质岩研究具有全球对比意义。并且,硅质岩的发育常与优质烃源岩伴生,如四川盆地五峰组–龙马溪组^[12],华南地区茅口组^[18–19],准噶尔盆地风城组^[8,11,13]和芦草沟组^[10],以及三塘湖盆地芦草沟组^[20]等。事实上,硅质岩发育对有机质富集^[21]、页岩油气储集^[22]和开发^[14–15]等方面都具有重要意义,贯穿了页岩油“生烃–成藏–开发”的全部过程。

准噶尔盆地芦草沟组沉积于咸水湖相环境中^[23–24],其中硅质岩发育,为陆相页岩中硅质岩成因机制及油气意义研究提供了基础。因此,在对博格达山北麓芦草沟组硅质岩矿物组成、沉积特征、成因机制等方面研究的基础上,进一步分析了硅质岩发育对陆相页岩有机质富集、储集空间形成和页岩油开发等方面的影响。

1 地质背景

准噶尔盆地位于新疆北部(图1a),是中国重要的大型含油气盆地,面积约为 $13 \times 10^4 \text{ km}^2$ ^[25]。盆地西北以哈拉阿拉特山和扎伊尔山为界,东北以阿尔泰山和克拉美丽山为界,南以伊林黑比尔根山和博格达山为界。其中,芦草沟组主要分布于准噶尔盆地东南部的博格达山周缘地区(图1b)。博格达地区主要经历了3个构造和沉积演化阶段:石炭纪–中二叠世为裂谷盆地的主要伸展期,沉积环境由滨浅海逐渐过渡为湖相^[26];晚二叠世–早三叠世博格达地区仍然为伸展盆地,局部地区开始隆升,沉积环境主要为冲积扇和三角洲^[27];三叠纪–白垩纪由于印支运动导致博格达山大规模隆升,从而分隔了准噶尔盆地和吐哈盆地^[28]。

中二叠统芦草沟组为裂谷盆地中的一套湖相细粒沉积岩,沉积期伴有间歇性的火山–热液活动^[29],使芦草沟组中火山碎屑和热水沉积物广泛发育,如凝灰岩、热液白云岩、热液黄铁矿和硅质岩等^[6,10,30]。芦草沟组沉积厚度为200~1 000 m^[4],古气候为温暖湿润气候^[24],水体环境以还原性较强^[31]的咸水–微咸水^[23]为主,在沉积期具有较高的古生产力^[29],富含藻类、古鲕鱼和介形虫化石^[32]。特殊的沉积环境造成了芦草沟组混合沉积的特征,岩性主要为富有机质页岩和暗色泥岩,夹有粉砂岩和碳酸盐岩薄层,并且在纵向上广泛分布着结核状硅质岩(图2)^[10]。

2 样品与方法

对博格达山XLK1井芦草沟组岩心中硅质岩发育层段进行了系统取样,所有样品均用金刚石线锯切为镜像对称的两块,其中一块用于制作薄片,另一块进行抛光用于元素面扫和粉末、小块样品取样。用牙钻对硅质岩样品和围岩样品进行微区取样,取样后的粉末

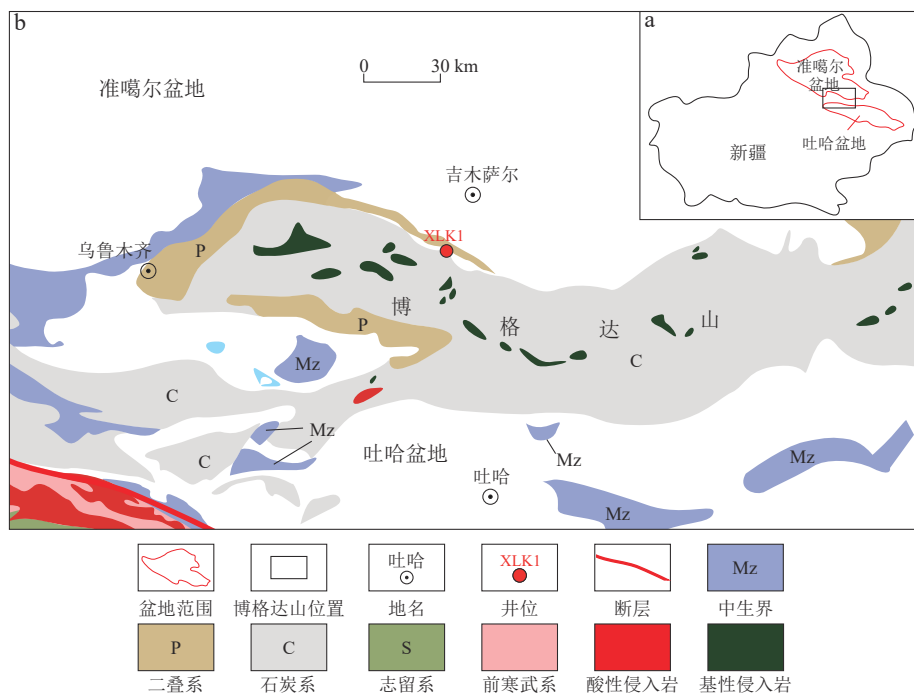
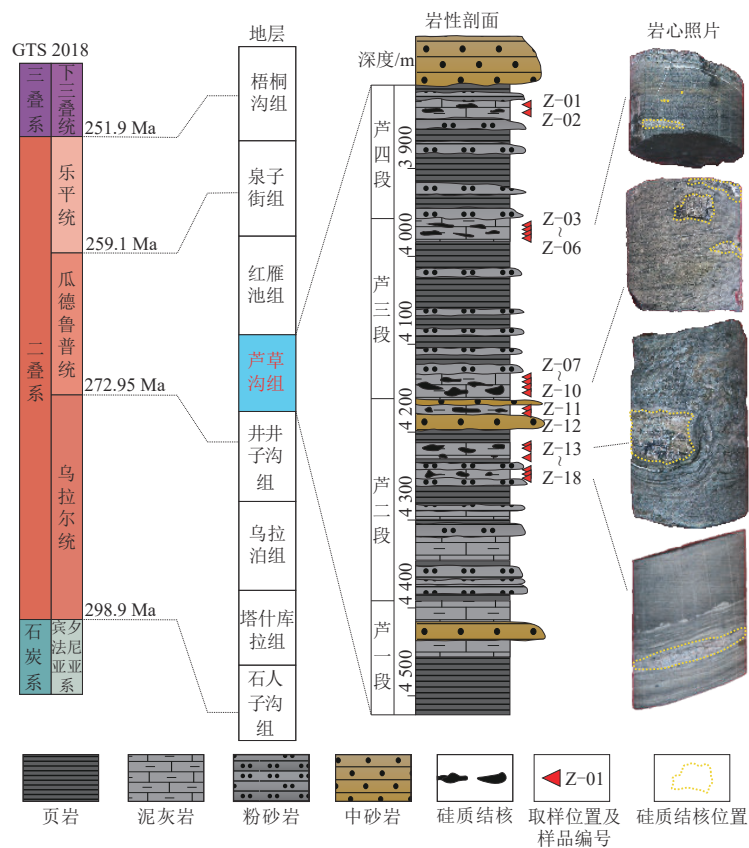


图1 准噶尔盆地博格达山周缘地区地质简图及XLK1井位置

Fig. 1 Geological sketch showing the surrounding area of Bogda Mountain in the Junggar Basin with the location of Well XLK1

a. 盆地位置示意图; b. 博格达山周缘地质图

图2 准噶尔盆地东南缘地层年龄格架^[25]及XLK1井芦草沟组综合柱状图^[10]Fig. 2 Ages of the Permian strata in the southeastern Junggar Basin^[25] and lithology column of the Lucaogou Formation in Well XLK1^[10]

进一步用研钵研磨至小于200目,用于X射线衍射(XRD)分析、主微量元素测试、岩石热解和总有机碳含量(TOC)分析。

用布鲁克公司的微区X射线荧光分析仪(型号为Tornado M4)对抛光后的岩石样品进行了元素平面扫描。元素平面扫描的点间距为30 μm ,每个点的测量时间为7 ms,测试方法参照刘可禹和刘畅(2019)^[33]。对18个硅质岩粉末样品进行了XRD分析,测试条件为Cu靶,测试电压为40 kV,测试功率为12 kW。用ICP-OES和ICP-MS分别对主元素和微量元素进行了测试。对共计24个泥页岩样品进行了岩石热解分析。对9个泥页岩样品进行了镜质体反射率(R_o)测试, R_o 测试和前处理参照行业标准《SY/T5124—2012沉积岩中镜质体反射率测定方法》^[34]。TOC测试在中科院地质与地球物理研究所进行,测试仪器为高频红外碳硫分析仪。

3 结果与讨论

3.1 硅质岩发育特征及成因

芦草沟组硅质岩在纵向上分布较为广泛,以结核状为主,主要附存于泥灰岩中,在多层中叠置出现(图2)。横向上硅质结核顺层发育,长轴与层面平行(图3a)。结核中软沉积物变形构造非常发育,并且表现为与围岩一致的变形特征(图3a,b),说明硅质结核形成于围岩尚未固结时。XRD分析表明,硅质结核主要由石英组成,含量为82.10%~98.90%,平均为92.38%,含有少量方解石,含量为0~17.9%,平均为7.22%,并且在部分样品中还检测到了少量黏土矿物,平均含量为0.41%(图4)。

硅质结核中石英在单偏光下呈透明状且表面干

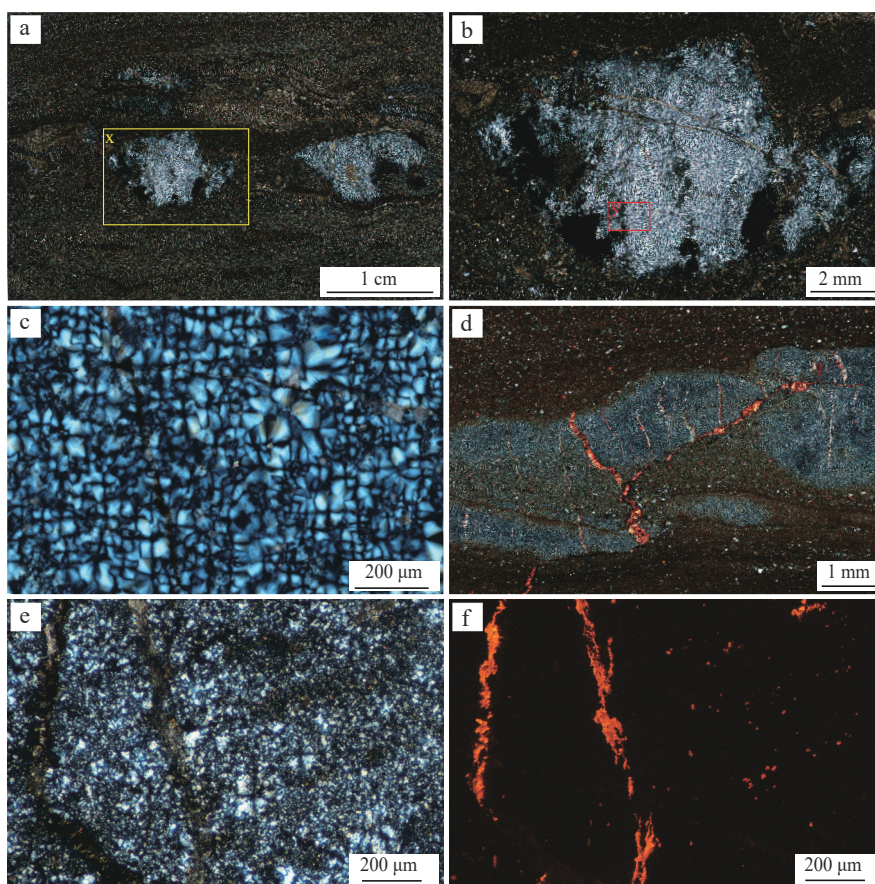


图3 准噶尔盆地XLK1井芦草沟组硅质结核岩石学特征显微照片

Fig. 3 Petrological characteristics of siliceous nodules in the Lucaogou Formation in Well XLK1, Junggar Basin

a. 硅质岩呈结核状顺层分布,软沉积物变形构造发育,埋深4 224.91 m,正交光;b. 硅质岩中裂缝发育,正交光,对应图a中x区域(黄框);c. 硅质岩主要由隐-微晶石英组成,正交光,对应图b中y区域(红框);d. 硅质岩中裂缝发育,被方解石完全充填,茜素红染色,埋深4 202.25 m,正交光;e. 硅质结核内部裂缝发育,被方解石完全充填,埋深4 226.72 m,正交光;f. 图e对应的阴极发光照片,隐-微晶石英不发光,方解石脉发橙色光

净,在正交偏光下呈亚矩形格子状,结晶程度较差,主要为隐-微晶结构(图3b—e),说明石英以自生为主,受成岩作用影响较小。阴极发光进一步验证了石英的

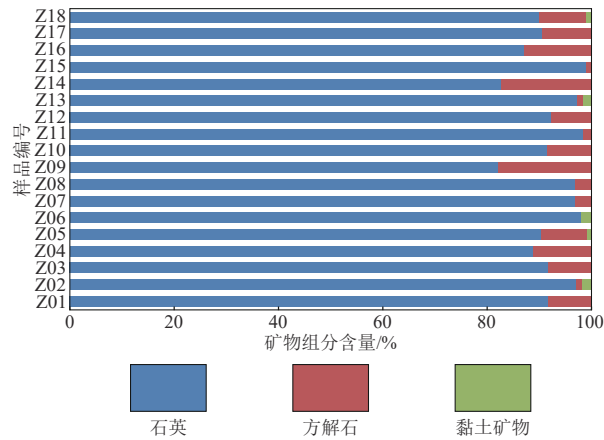


图4 准噶尔盆地芦草沟组硅质岩矿物组成

Fig. 4 Mineral composition of siliceous rocks in the Lucaogou Formation, Junggar Basin

自生成因,通常发蓝紫色光的石英源自于深成岩或火山岩,发红棕色光的石英源自于区域变质岩,无阴极发光的石英多为自生成因^[14]。芦草沟组硅质岩中石英均无阴极发光(图3f),结合硅质岩中软沉积物变形构造广泛发育等特征,说明硅质岩为同沉积期形成。硅质结核中方解石主要为裂缝充填物,茜素红染色后呈红色(图3d),为成岩期胶结作用的产物。

扫描电镜观察表明,硅质结核中发育重晶石、锰菱铁矿、镍铁矿、蓝铜矿和黄铁矿等自生含金属矿物,在背散射下呈灰白色,自形程度较高,为半自形-自形结构(图5)。重晶石(图5a)和镍铁矿(图5b)主要呈粒状附存于隐-微晶石英中,平均粒径分别为5 μm 和33 μm 。锰菱铁矿(图5c)和蓝铜矿(图5d)主要为片状集合体状分布于硅质结核内部,平均粒径分别为3 μm 和17 μm 。黄铁矿呈微晶集合体状或分散状,平均粒径为5 μm (图5e,f)。硅质岩结构致密,石英晶间孔径大多小于3 μm ,小于自生含金属矿物平均粒径,且在

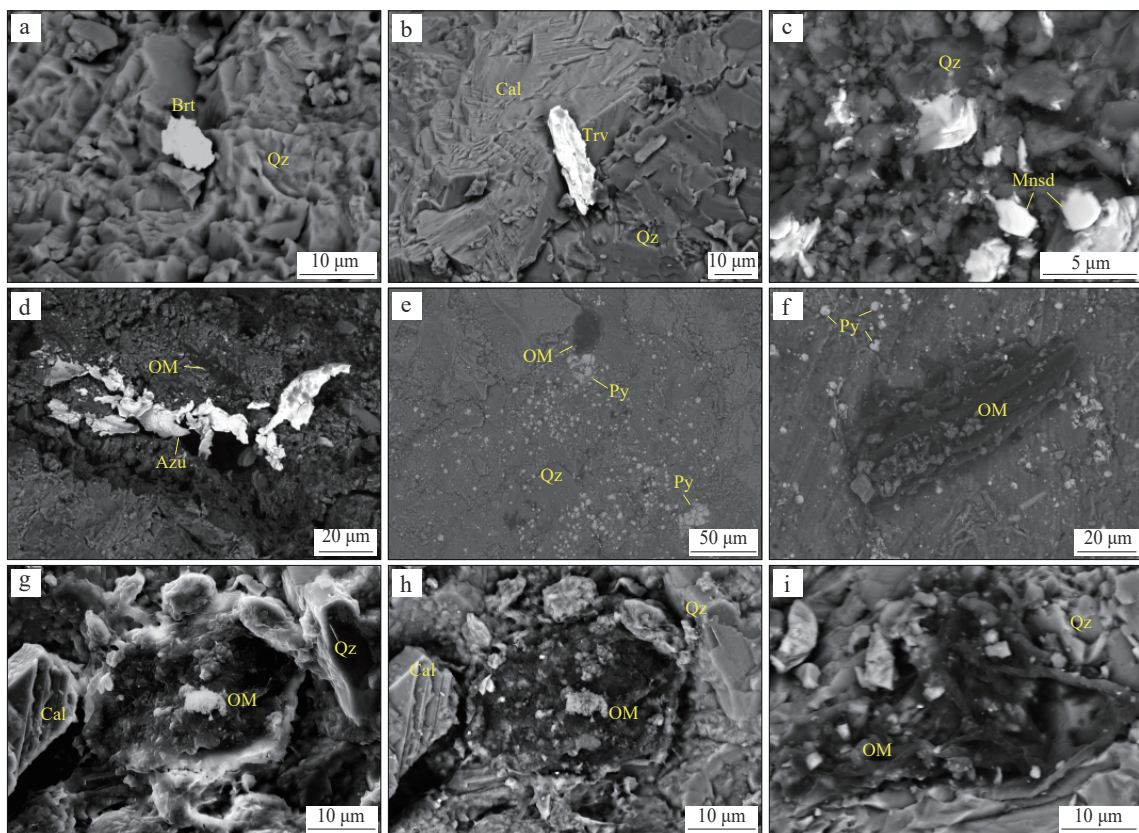


图5 准噶尔盆地 XLK1 井芦草沟组硅质岩及围岩扫描电镜特征照片

Fig. 5 Scanning electron microscope characteristics of siliceous rocks and surrounding rocks in the Lucaogou Formation in Well XLK1, Junggar Basin
a. 硅质岩中自生重晶石发育,埋深4 226.72 m,背散射;b. 硅质岩中的方解石及微晶镍铁矿,埋深4 226.73 m,背散射;c. 硅质岩中锰菱铁矿发育,呈微晶分散状,埋深4 226.90 m,背散射^[10];d. 硅质岩中发育蓝铜矿及少量有机质,埋深4 226.90 m,背散射;e. 硅质岩顶部黄铁矿及有机质发育,埋深4 226.91 m,背散射;f. 硅质岩顶部的有机质及黄铁矿,埋深4 226.91 m,背散射;g. 硅质岩上覆围岩中的有机质,二次电子;h. 图g对应的背散射图像;i. 硅质岩上覆围岩中的有机质,埋深4 226.72 m,背散射

Brt. 重晶石;Qz. 石英;Cal. 方解石;Trv. 镍铁矿;Mnsd. 锰菱铁矿;Azu. 蓝铜矿;OM. 有机质;Py. 黄铁矿

硅质岩内部裂缝中没有重晶石、锰菱铁矿、镍铁矿和蓝铜矿的发育,说明这些含金属矿物与硅质结核应为同期形成。此外,在硅质结核顶部还发现了分散状有机质,直径为 $4 \sim 43 \mu\text{m}$,在背散射下呈深灰色(图5e,f)。

硅质岩按照硅质来源可分为陆源硅^[35]、生物硅^[36]和深源硅^[37]3类,其中陆源硅主要源自于源区铝硅酸岩风化^[38],由河流携带至汇水区中^[17];生物硅指硅藻、放射虫和海绵等生物在生物硅化作用下形成的硅,典型特征为硅质岩中富含硅质生物化石^[39];深源硅指由火山或热液等深部流体活动带来的硅^[37],其中热液成因硅常伴生重晶石、含锰矿物、含镍矿物以及含铜矿物^[40-41]。芦草沟组硅质结核中大量分散状的自生含金属自生矿物表明硅源可能为热液流体。

硅质岩元素特征验证了硅质岩的热液成因(图6,

图7)。热液流体通常具有较高的Fe元素和Mn元素含量^[9,37],而Al元素含量通常与陆源输入强度有关^[42]。与围岩相比,硅质结核具有高Si元素而贫其他元素的特征,在硅质结核发育层的围岩中Al元素含量相对较低(图6c),而Fe元素(图6e)和Mn元素含量相对较高(图6f)。稀土元素和Y元素含量是硅质岩硅源判识的常用参数^[19],芦草沟组硅质岩稀土元素具有总含量低(平均为 $2.85 \mu\text{g/g}$)、重稀土元素相对富集(Nd_N/Yb_N 比平均为0.13)、低Y/Ho比(平均为28.15)和显著正Eu异常(δEu 值平均为2.78)等与热液流体^[43]和热液硅质岩^[10]一致的特征(图7),说明XLK1井芦草沟组硅质岩为热液成因,并且每一层硅质结核都记录了一次热液流体活动脉冲,揭示了博格达裂谷在中二叠世存在间歇性热液活动。

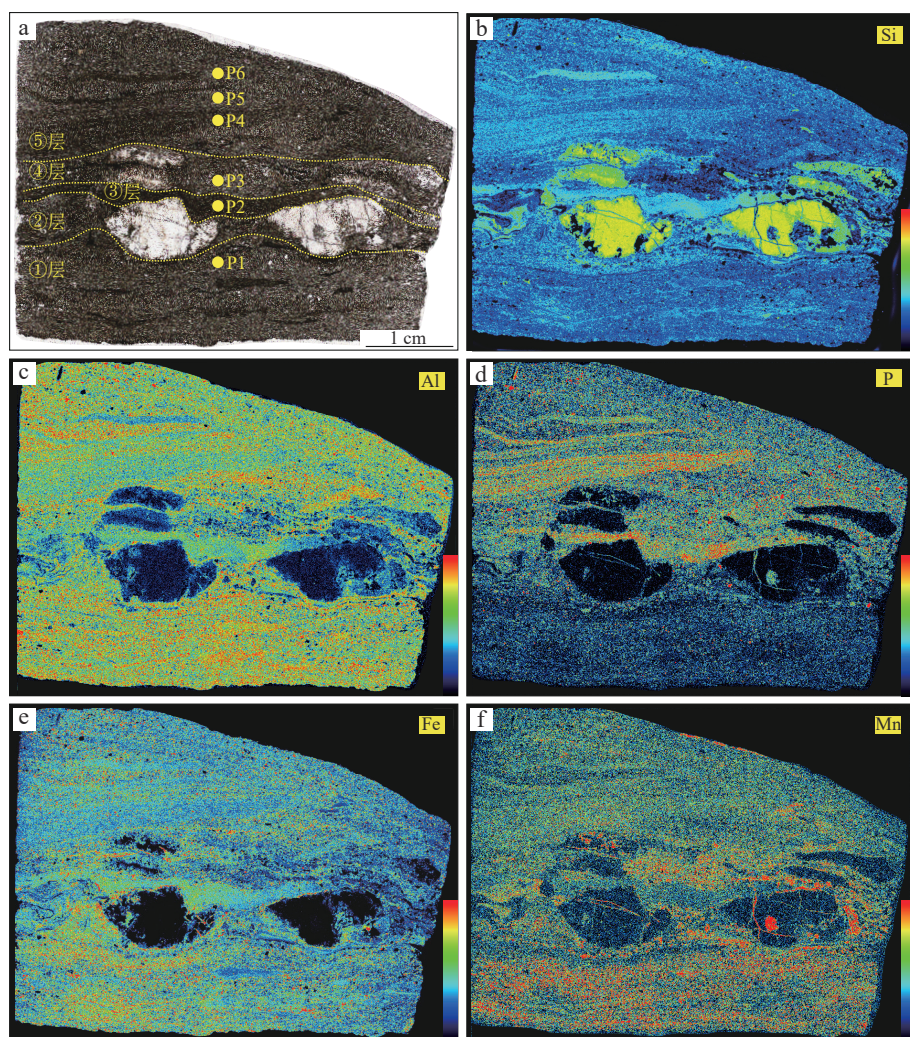


图6 准噶尔盆地XLK1井芦草沟组硅质岩(埋深4 224.91 m)元素面扫结果

Fig. 6 The Micro-XRF mapping results of siliceous rock in the Lucaogou Formation in Well XLK1, Junggar Basin

a. 硅质结核样品显微照片,单偏光;b. Si-K α ;c. Al-K α ;d. P-K α ;e. Fe-K α ;f. Mn-K α

(①层—⑤层代表泥灰岩分层;P1—P6代表取样位置;色标由蓝色到红色表示元素强度逐渐增大。)

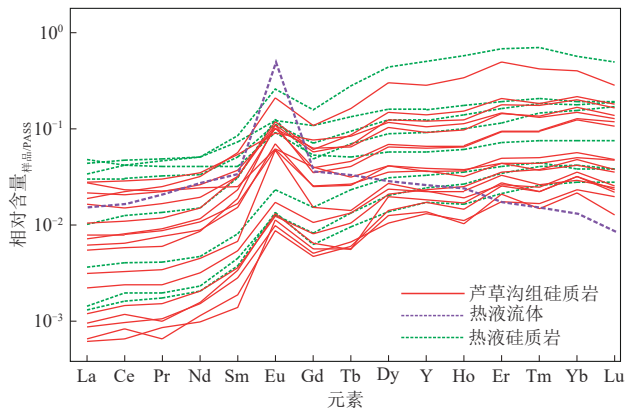


图7 澳大利亚后太古宙平均页岩(PAAS)标准化的准噶尔盆地芦草沟组硅质结核、热液流体^[43]和热液硅质岩^[10]稀土元素和钇(Y)元素配分曲线

Fig. 7 PAAS-normalized rare earth element and yttrium values of siliceous nodules, hydrothermal fluid^[43] and hydrothermal siliceous rocks in the Lucaogou Formation^[10]

3.2 硅质岩对有机质富集的影响

在海相环境中,硅质岩的发育常与古生产力勃发相伴生,这是由于硅是放射虫和海绵骨针等硅质生物生长所必需的营养元素,因此硅元素含量可作为海洋古生产力大小的有效替代指标^[21,44]。不同于海相环境,芦草沟组硅质岩形成于咸水湖相环境中,虽然缺少硅质生物发育,但是硅质结核伴生的热液活动同样对有机质富集产生了重要影响^[29]。通常热液活动对有机质富集主要有两类影响:①热液活动带来的营养物质使藻类勃发,进而提高古生产力^[45];②热液流体带来的大量金属离子会使湖水盐度增加,使水体更易发生盐度分层,为沉积有机质的保存创造良好条件^[46]。

通过元素面扫对硅质结核发育样品的矿物组成和地球化学特征进行了分析,元素面扫结果中从蓝色到红色表示元素强度逐渐增大,与元素含量呈正相关关系^[33]。基于元素面扫结果,将发育硅质结核的泥灰岩划分为正常沉积围岩、热液活动层和热液活动后围岩,并进一步对古生产力和古水体环境进行了示踪。如图所示,①层为正常沉积围岩,形成于热液活动前;②层和④层为热液活动层,③层和⑤层为热液活动后围岩(图6)。正常沉积围岩中TOC仅为0.3%,而热液活动后围岩的TOC增加到1.5%,并且在硅质结核的顶部及热液活动后围岩中均发现了大量分散有机质(图5g—i),说明热液活动对有机质富集起到了建设性作用。

通常可用Al元素含量指示陆源碎屑输入,随陆源输入增加,Al元素含量逐渐增加^[9,37],而Fe元素和Mn元素的含量可指示热液活动^[47]。为了消除陆源输入

的影响,用Fe/Al和Mn/Al来评价热液活动的影响。P元素、K元素和Ba元素是生物发育所必需的重要营养元素^[48-49],可用P/Al、K/Al和Ba/Al反映古生产力强度^[50]。正常沉积围岩表现为高Al元素含量(平均为11.247%),低P/Al(平均为0.003),K/Al(平均为0.356),Ba/Al(平均为0.019),Fe/Al(平均为0.866)和Mn/Al比值(平均为0.034)的特征(图8),说明正常沉积层主要受陆源输入影响,古生产力相对较低。热液活动层具有低Al元素含量(平均为6.842%),中等P/Al(平均为0.022)和K/Al比值(平均为0.176),高Ba/Al(平均为0.049),Fe/Al(平均为1.349)和Mn/Al比值(平均为0.101)的特征,说明热液活动层受陆源输入影响相对较小,而受热液活动影响较大(图8)。热液活动后围岩具有Al元素含量相对较高(平均为7.259%),高P/Al(平均为0.126)和K/Al比值(平均为0.414),低Ba/Al(平均为0.011),Fe/Al(平均为1.262)和Mn/Al比值(平均为0.043)的特征(图8)。P/Al和K/Al较正常沉积围岩的显著升高说明在热液活动后,古生产力显著提高。然而,Ba/Al却表现出与P/Al和K/Al相反的特征,这是由于海洋/湖泊中钡的来源主要有3种,分别为陆源钡、生物钡和热液钡,其中仅生物钡能指示古生产力^[51]。Ba/Al比值与Fe/Al和Mn/Al比值的变化趋势较为一致,说明在芦草沟组含硅质结核样品中Ba元素含量受到了

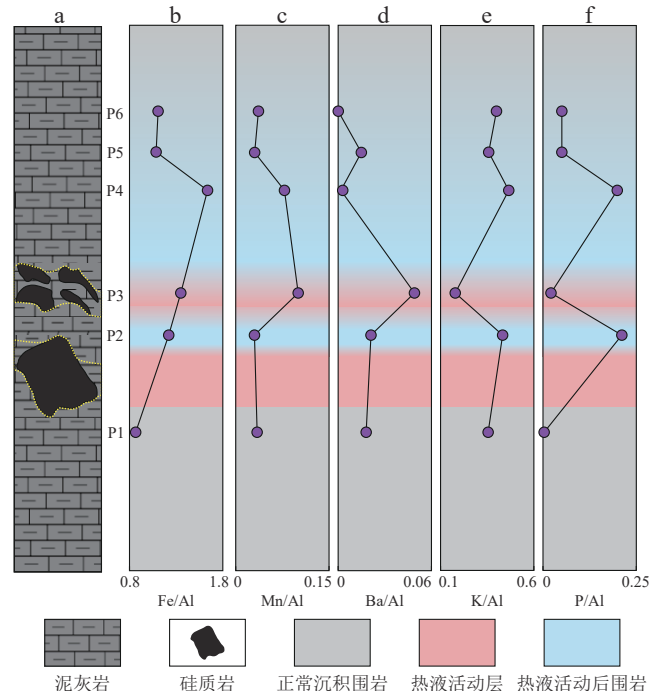


图8 准噶尔盆地芦草沟组含硅质岩样品元素含量比值垂向变化
Fig. 8 Vertical variations in element ratios of siliceous rock samples in the Lucaogou Formation, Junggar Basin

a. 取样位置(P1—P6位置见图6a);b—f. 元素含量比值垂向变化

热液流体影响,因此不适用作为古生产力替代指标。

由于元素面扫测量的元素有限,进一步通过 ICP-MS 和 ICP-OES 测量的主、微量元素含量来分析和对比热液活动前、后的水体环境变化。Sr 元素含量、Sr/Ba 比和 B/Ga 比能有效反映古盐度,随古盐度增加逐渐增加^[52]。正常沉积围岩的 Sr 元素含量(为 888 $\mu\text{g/g}$)、Sr/Ba(为 0.86)和 B/Ga 比值(为 3.57)均分别低于热液活动后围岩的 Sr 元素含量(为 1 182 $\mu\text{g/g}$)、Sr/Ba(为 0.94)和 B/Ga 比值(为 5.85),说明热液活动使湖泊古盐度增加。U/Th、V/Cr 和 Ni/Co 比是指示水体氧化-还原环境的有效参数^[53],与水体还原性呈正相关关系。和正常沉积围岩(U/Th=0.40, V/Cr=1.54 和 Ni/Co=23.10)相比,热液活动后围岩的 U/Th(为 1.17),V/Cr(为 2.19)和 Ni/Co 比值(为 42.67)均明显增大,说明热液活动后,水体还原性增强。

热液活动前后古生产力和古水体环境指标的变化综合说明芦苇沟组中硅质结核伴生的热液活动带来了 Fe、Mn、Cu 和 Ni 等营养元素,促使古生产力勃发,并且热液流体携带的大量金属阳离子使湖水盐度增加,促使湖水盐度分层,进一步导致底水还原性增强。热液活动后,古生产力逐渐降低,说明热液活动对古生产力的影响具有一定持续性,热液活动停止后,随着湖泊水体养分的逐渐消耗,古生产力逐渐降低。

3.3 硅质岩对储集空间的影响

硅质岩的发育有利于储集空间的发育和保持^[44]。芦苇沟组页岩油储层具有孔径小、渗透率差异大、储层非均质性强等特点,寻找页岩中的“甜点”是页岩油勘探的重点方向^[54]。与海相硅质岩中孔隙以生物体腔

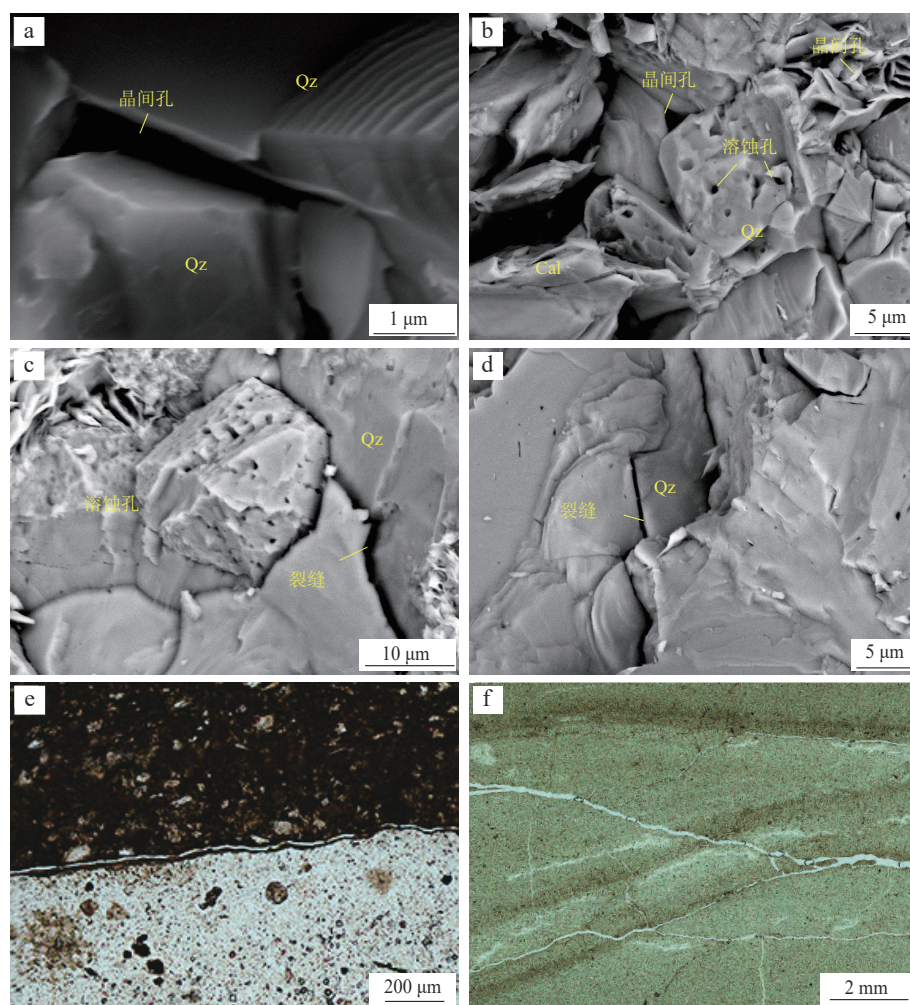


图9 准噶尔盆地 XLK1 井芦苇沟组硅质岩储集空间类型显微照片

Fig. 9 Microscopic images showing types of reservoir space of siliceous rocks in the Lucaogou Formation in Well XLK1, Junggar Basin
a. 硅质岩石英晶间孔发育,埋深4 224.92 m,二次电子SEM;b. 石英晶间孔及内部溶蚀孔,黏土矿物晶间孔,裂缝被方解石完全充填,埋深4 224.91 m,背散射SEM;c. 硅质岩中石英溶蚀孔发育,石英被溶蚀成海绵状,张裂缝发育,埋深4 226.72 m,背散射SEM;d. 硅质岩中张裂缝发育,石英断口呈贝壳状,埋深4 226.90 m,背散射SEM;e. 硅质岩结核边缘发育张裂缝,埋深4 202.25 m,单偏光;f. 硅质岩内部张裂缝发育,埋深4 221.75 m,单偏光
Qz. 石英;Cal. 方解石

孔为主不同^[21],陆相硅质岩的储集空间类型更加复杂。芦草沟组硅质岩中孔隙和裂缝非常发育,储集空间类型多样,晶间孔、溶蚀孔和裂缝均有发育(图9)。

晶间孔主要发育在微晶石英晶体间(图9a,b),孔径为0.66~3.51 μm ,平均为2.05 μm 。溶孔主要发育在石英晶体表面(图9b,c),孔径为0.20~1.26 μm ,平均为0.55 μm 。部分石英晶体经历较强溶蚀作用,被溶蚀为“海绵”状(图9c)。裂缝主要分布于硅质岩内部和硅质结核边缘(图9d—f),按裂缝开度可分为大缝(>100 μm)、中缝(10~100 μm)、小缝(1~10 μm)和微缝(<1 μm)。扫描电镜下,可见硅质岩内部微缝十分发育,开度为0.13~0.84 μm ,平均为0.52 μm 。显微镜下可见小缝、中缝和大缝的发育,大缝的开度为122.63~175.19 μm ,平均为160.85 μm ,中缝的开度为22.30~65.30 μm ,平均为44.59 μm ,小缝的开度

为7.82~9.72 μm ,平均为8.49 μm 。硅质岩内部裂缝常被方解石完全充填,部分裂缝未充填,未充填缝可能形成于方解石胶结作用之后。虽然硅质岩中孔隙的孔径较小,以微孔为主,孔隙连通性较差,但是硅质岩中广泛发育的裂缝有效沟通了孔隙,为页岩油储集提供了空间。此外,由于芦草沟组硅质岩形成于大规模压实作用前,有助于提高岩石抗压实作用能力,从而对储层起到较好的保护作用。

3.4 硅质岩对页岩油开发的影响

硅质岩的发育有利于储层压裂改造和页岩油开发^[21]。中国陆相页岩油具有油质重、黏度高、流动能力差、黏土矿物含量高和脆性较低等特点,储层非均质性强、有机质成熟度低和脆性矿物含量低是制约页岩油开发的重要因素^[55-56]。脆性矿物含量越

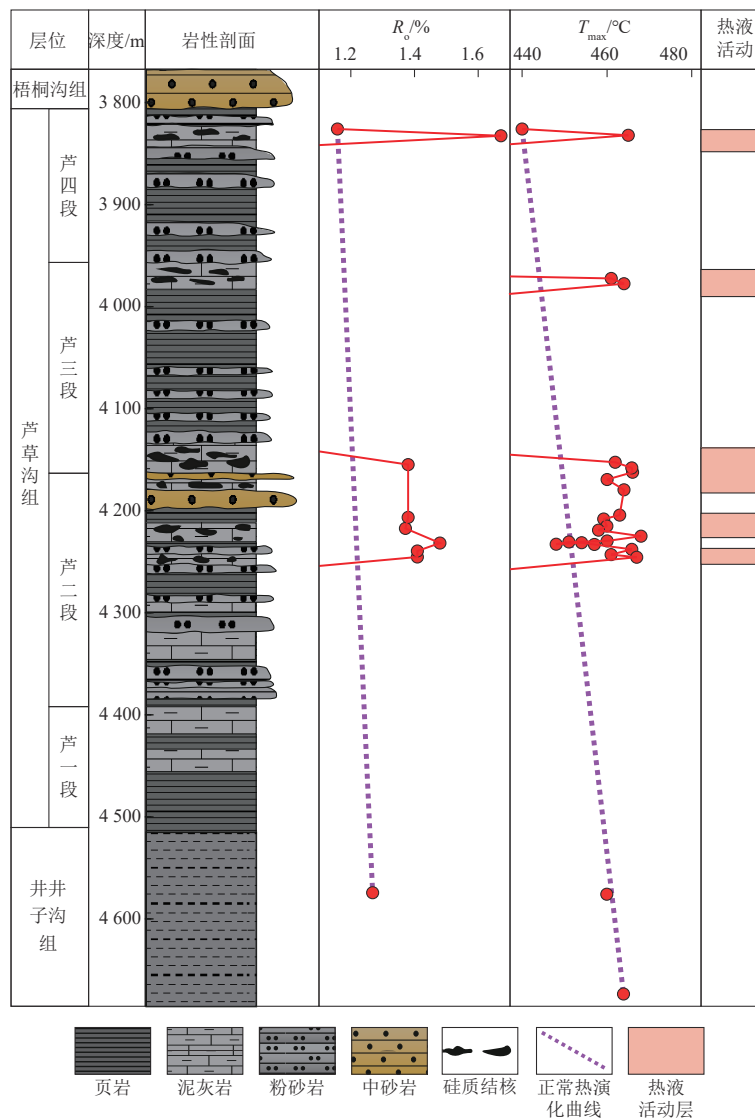


图10 准噶尔盆地XLK1井芦草沟组成熟度参数纵向变化

Fig. 10 Vertical variation of maturity of the Lucaogou Formation in Well XLK1, Junggar Basin

高、储层可压裂性就越强^[44],石英属于脆性矿物,硅质岩的发育有助于提高芦草沟组储层脆性矿物含量。芦草沟组硅质岩中,石英含量平均为92.38%,且芦草沟组硅质岩通常顺层发育,改善了储层可压裂性。此外,深源物质参与有助于有机质生烃演化^[30],XLK1井芦草沟组 R_o 和最高热解峰温(T_{max})正异常层段与硅质岩发育层段具有较好的对应关系(图10),说明硅质岩伴生的热液流体活动参与了有机质生烃演化,从而使页岩油黏度降低,可动性增强。

4 结论

1) 芦草沟组硅质岩主要呈结核状顺层发育,软沉积物变形广泛发育,结核主要由隐-微晶石英组成,内部可见重晶石、锰菱铁矿、镍铁矿和蓝铜矿等自生含金属矿物,稀土元素具有总量低,重稀土元素相对富集和正Eu异常等特征,为典型同沉积热液成因硅质岩,说明芦草沟组沉积期博格达裂谷盆地存在间歇性热液活动。

2) 硅质岩伴生的热液流体携带了大量Fe、Mn、Cu和Ni等营养元素,有助于藻类勃发,提高古生产力,并且热液中的金属阳离子使湖水盐度增加,促使湖水分层,进一步导致底水还原性增强,为有机质保存提供了良好条件。

3) 硅质岩中晶间孔、溶蚀孔和裂缝的发育有利于页岩油储层的发育和沟通。并且,硅质岩的发育提高了储层中石英等脆性矿物含量,使储层可压裂性增强,硅质岩伴生的热液活动参与了有机质热演化,提高了页岩油可动性,有助于页岩油开发。

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