

鄂尔多斯盆地氦源岩特征及生氦潜力

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摘要:氦源是氦气资源形成的首要条件,盆地基底变质岩、泥岩、泥质白云岩、煤和铝土岩等沉积岩为潜在氦源岩。通过野外地质调查、重磁资料解释、岩心描述与主、微量元素测试等方法,研究了鄂尔多斯盆地2大类5套潜在氦源岩及其成氦潜力。研究结果表明:鄂尔多斯盆地的潜在氦源岩分为基底型和沉积型两大类。基底型氦源岩发育在太古宇陆块及其上叠加的古元古界之中,岩性主要为高级变质片麻岩-变粒岩、大理岩、混合岩和花岗片麻岩,U和Th元素平均丰度分别为 3.15×10^{-6} 和 12.38×10^{-6} ,生氦强度为 $0.735 \times 10^{-6} \text{ cm}^3/(\text{a} \cdot \text{g})$ 。沉积型氦源岩主要发育在中元古界长城系沉积变质岩和古生界沉积岩之中。长城系黑色板岩主要分布在盆地北部和西南部,U和Th平均丰度分别为 2.36×10^{-6} 和 8.28×10^{-6} ,生氦强度为 $0.522 \times 10^{-6} \text{ cm}^3/(\text{a} \cdot \text{g})$;下古生界下奥陶统马家沟组泥质白云岩分布在盆地中部及东部,U和Th平均丰度分别为 1.71×10^{-6} 和 9.80×10^{-6} ,生氦强度为 $0.487 \times 10^{-6} \text{ cm}^3/(\text{a} \cdot \text{g})$;上古生界石炭系-二叠系泥岩和煤在全盆地广泛分布,太原组泥岩U和Th平均丰度分别为 9.69×10^{-6} 和 22.68×10^{-6} ,生氦强度为 $1.82 \times 10^{-6} \text{ cm}^3/(\text{a} \cdot \text{g})$,太原组煤U和Th平均丰度分别为 16.12×10^{-6} 和 44.13×10^{-6} ,生氦强度为 $3.21 \times 10^{-6} \text{ cm}^3/(\text{a} \cdot \text{g})$;上古生界石炭系铝土岩主要分布在盆地东部和西南部,U和Th平均丰度分别为 7.14×10^{-6} 和 38.57×10^{-6} ,生氦强度为 $1.97 \times 10^{-6} \text{ cm}^3/(\text{a} \cdot \text{g})$ 。鄂尔多斯盆地西南部发育各类氦源岩,具有“多源供氦”特征。该研究为鄂尔多斯盆地氦气资源勘查奠定了基础。

关键词:生氦潜力;氦源岩;氦气;天然气藏;鄂尔多斯盆地

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Characteristics and helium-generating potential of helium source rocks in the Ordos Basin

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Abstract: Helium sources act as the primary element in generating helium resources. Potential helium source rocks encompass basement metamorphic rocks and various sedimentary rocks, such as mudstone, argillaceous dolomite, coal, and bauxite, in basins. Based on field geological surveys, gravity-magnetic data interpretation, core description, and major- and trace-element analyses, we investigate five suites of potential helium source rocks of two categories in the Ordos Basin and their helium-generating potential. The results show that two categories of potential helium source rocks, namely basement metamorphic rocks and sedimentary rocks, are developed in the studied basin. The helium source rocks of the basement type are found in the Archean land block and the superimposed Paleoproterozoic strata, dominated by high-grade metamorphic gneiss, granulite, marble, migmatite, and granitic gneiss, with average abundances of U and Th of 3.15×10^{-6} and 12.38×10^{-6} , respectively, and a helium-generating intensity of $0.735 \times 10^{-6} \text{ cm}^3/(\text{a} \cdot \text{g})$. On the other hand, the helium source rocks of the sedimentary type are primarily found in Changchengian metasedimentary rocks of the Mesoproterozoic strata and the Paleozoic sedimentary rocks. The Changchengian black slate is widely seen

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in the northern and southwestern Ordos Basin, and exhibits average abundances of U and Th of 2.36×10^{-6} and 8.28×10^{-6} , respectively, and a helium-generating intensity of $0.522 \times 10^{-6} \text{ cm}^3/(\text{a} \cdot \text{g})$. The argillaceous dolomite in the Lower Ordovician Majiagou Formation of the Lower Paleozoic strata, found in the central and eastern Ordos Basin, displays average abundances of U and Th of 1.71×10^{-6} and 9.80×10^{-6} , respectively, and a helium-generating intensity of $0.487 \times 10^{-6} \text{ cm}^3/(\text{a} \cdot \text{g})$. The Carboniferous-Permian mudstones and coals of the Upper Paleozoic strata are extensively distributed throughout the Ordos Basin. The Taiyuan Formation mudstones, among others, show average abundances of U and Th of 9.69×10^{-6} and 22.68×10^{-6} , respectively, and a helium-generating intensity of $1.82 \times 10^{-6} \text{ cm}^3/(\text{a} \cdot \text{g})$, while its coals exhibit average abundances of U and Th of 16.12×10^{-6} and 44.13×10^{-6} , respectively, and a helium-generating intensity of $3.21 \times 10^{-6} \text{ cm}^3/(\text{a} \cdot \text{g})$. The Carboniferous bauxite of the Upper Paleozoic strata, primarily occurring in the eastern and southwestern Ordos Basin, demonstrates average abundances of U and Th of 7.14×10^{-6} and 38.57×10^{-6} , respectively, and a helium-generating intensity of $1.97 \times 10^{-6} \text{ cm}^3/(\text{a} \cdot \text{g})$. Various helium source rocks are discovered in the southwestern Ordos Basin, suggesting a multi-source helium supply. Overall, this study lays the foundation for helium resource exploration in the Ordos Basin.

Key words: helium-generating potential, helium source rock, helium, natural gas reservoir, Ordos Basin

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作为现代高科技产业无可替代的关键元素,氦气是重要的战略资源,全球氦资源主要集中在美国、阿尔及利亚和俄罗斯等少数国家。中国氦资源对外依存度极高,进口氦气的渠道和来源地又极为集中和有限,严重制约中国高科技产业和国家安全^[1-3]。学者们对富氮常规气、致密气、水溶气藏形成条件开展了研究,取得的进展和认识为本文研究奠定了基础^[4-16]。中国已在中部四川、鄂尔多斯等盆地,西部塔里木、柴达木等盆地,东部松辽、海拉尔、苏北、渤海湾等盆地,南方三水等盆地中发现富氮、高氮天然气藏^[5-6, 16-25]。部分地区地热井和温泉中也有氦气发现^[26-27]。

天然气中的氦主要有大气源、壳源(岩石中放射性U和Th元素衰变形成)和幔源(地球形成时赋存于地球深部的原始氦),多数具有商业价值的氦气资源主要为壳源氦。壳源氦源岩又分为壳源远源型与壳源近源型两种类型^[28]。壳源远源型源岩主要包括高级变质片麻岩和变粒岩、大理岩、混合岩和花岗片麻岩等。例如渭河盆地富氮天然气藏的主要氦源岩为北秦岭广泛发育的富U花岗岩;东胜气田氦气主要来源于基底太古宇-元古宇的变质岩-花岗岩中U和Th元素的衰变^[25, 29]。壳源近源型源岩主要是富U和Th的沉积岩,特别是富有机质页岩。在地壳岩石中,由于吸附大量铀和钍元素且富有机质页岩的U和Th含量最高且氦气生成量最大,单位时间内富有机质页岩生成的氦气量是花岗岩的9~10倍。例如,四川盆地

五峰组-龙马溪组富有机质页岩的U和Th含量高、厚度大、分布广,氦气资源潜力大^[30-38]。

氦源是氦气资源形成的首要条件,氦源岩空间分布与氦源潜力评价是氦气资源勘查的关键问题。在野外地质调查、重磁资料解释、岩心描述及主、微量元素实验室分析测试等的基础上,系统分析不同U和Th元素含量岩石和矿物的成氮潜力,结合数值模拟定量评价鄂尔多斯盆地氦源潜力,厘定主力氦源;研究沉积盆地及其基底不同岩系的产氮丰度和强度。

1 氦气地质背景

鄂尔多斯盆地是发育在华北克拉通西部变质结晶基底之上的多旋回叠合盆地,依据盆地演化历史、构造特征及基底特性的不同,可将盆地划分成6个一级构造单元:伊盟隆起、晋西挠褶带、伊陕斜坡、渭北隆起、天环坳陷和西缘逆冲带^[39]。

盆地天然气资源丰富,目前已发现苏里格、榆林等多个气田。其中,东胜气田氦气平均含量为0.133%,苏里格气田东区、中区、西区和南区均为0.042%~0.108%,榆林气田为0.032%,神木气田为0.029%,子洲米脂气田为0.041%,黄龙气田为0.233%,庆阳气田为0.068%,表明鄂尔多斯盆地天然气具有较高的氦气含量(图1)。

鄂尔多斯盆地潜在氦源岩的岩石类型主要是太古

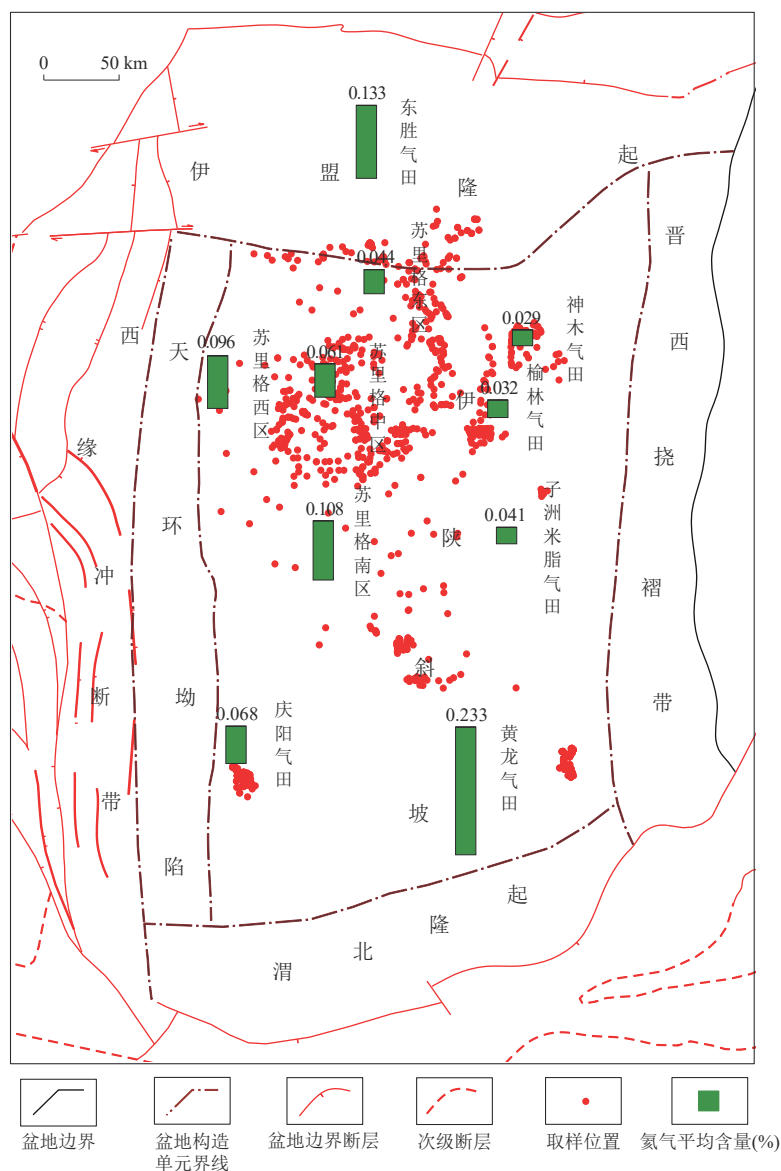


图1 鄂尔多斯盆地已发现含氦天然气田氦气丰度与样品采集点分布^[39]

Fig. 1 Map showing the abundance of helium and relevant sampling points in helium-bearing gas fields discovered in the Ordos Basin^[39]

宇-古元古界富含U和Th元素的花岗岩和花岗片麻岩等变质岩、中元古界长城系黑色板岩、下古生界下奥陶统马家沟组泥质含量较高的碳酸盐岩和上古生界上石炭统本溪组、下二叠统太原组与山西组的泥岩、煤及铝土岩(图2)。

2 基底型潜在氦源岩

基底经历了太古宙—古元古代构造运动,形成了一套整体磁性较强、以片岩、高级变质片麻岩和变粒岩为主的太古宇变质岩系与另一套整体磁性偏弱、以大理岩、混合岩和花岗片麻岩为主的古元古界变质岩系的稳定结晶基底(图3a—d)。太古宇陆块上叠加了古

元古界,从北到南依次为乌拉山群、上集宁群—贺兰山群、恒山—吕梁群、界河口群和涑水—中条群。

基底西部石榴夕线黑云斜长片麻岩与东部片麻状二云母花岗岩的钻孔样品进行SHRIMP锆石U—Pb定年,分别获得了 $2\,031\text{ Ma}\pm 10\text{ Ma}$ 和 $2\,035\text{ Ma}\pm 10\text{ Ma}$ 的年龄^[40]。北部乌拉山—大青山岩性主要为闪长质变质岩和碱性变质岩,狼山主要为闪长质变质岩,南部秦岭主要为花岗片麻岩。根据重力和磁力资料解释,太古宇和古元古界大致呈北东—南西向相间分布。根据放射性原理和铀、钍衰变方程[公式(1)]^[35],可计算出氦源岩的生氦强度。

$$^4\text{He} = 1.207 \times 10^{-13}U + 2.868 \times 10^{-14}Th \quad (1)$$

式中: ^4He 为氦源岩的生氦强度, $10^{-6}\text{ cm}^3/(\text{a}\cdot\text{g})$;U为铀

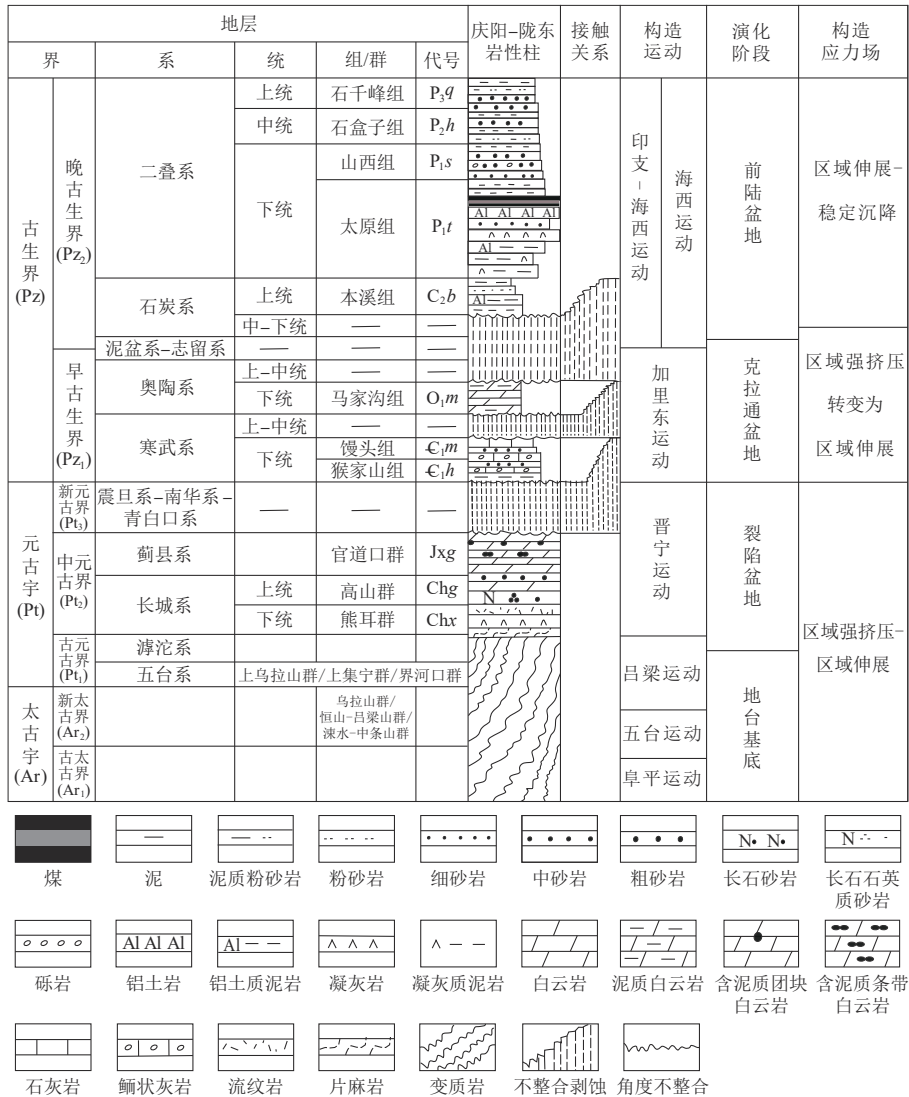


图2 鄂尔多斯盆地潜在氦源岩地层柱状图(据文献[39]修改)

Fig. 2 Stratigraphic column of potential helium source rocks in the Ordos Basin (modified from reference [39])

丰度, 10^{-6} ; Th 为钍丰度, 10^{-6} 。

盆地基底岩石 U 和 Th 平均丰度分别为 1.29×10^{-6} 和 8.19×10^{-6} , 生氮强度为 $0.391 \times 10^{-6} \text{ cm}^3/(\text{a} \cdot \text{g})$ (表1)。鄂尔多斯地区基底厚度远大于上覆沉积层, 具有较高的生氮潜力。

3 沉积型潜在氦源岩

基底形成之后, 盆地经历了晋宁、加里东、海西、印支、燕山及喜马拉雅等多期构造运动, 形成了以陆相碎屑岩和碳酸盐岩为主沉积盖层, 厚约 $3000 \sim 8000 \text{ m}$ [41-48]。

3.1 长城系

中元古界长城系具明显裂陷槽式(拗拉谷)沉积特征, 盆地西缘及南缘为主裂陷槽, 可能代表了秦、祁、贺三

又裂谷系最早活动的沉积记录, 厚度多在千余米及以上, 向盆地西缘深拗陷区有明显加厚趋势, 沉积厚度局部可达 $2000 \sim 3000 \text{ m}$ 。除了盆地东部, 长城系在盆地均有分布, 最大厚度可达 1000 m , 岩性主要为石英砂岩(图3e)、长石(岩屑)石英砂岩、炭质板岩及结晶灰岩(大理岩)。长城系黑色板岩是潜在氦源岩, 主要分布在盆地北部和西南部。中元古代初始年龄取 1717 Ma , U 和 Th 元素平均丰度为 1.29×10^{-6} 和 2.49×10^{-6} , 生氮强度为 $0.227 \times 10^{-6} \text{ cm}^3/(\text{a} \cdot \text{g})$ 。由于中元古界长城系潜在氦源岩岩石年代较老, U 、 Th 元素丰度较高, 在平面和深度的发育规模也较大, 可见其具有较高的生氮潜力。

3.2 下奥陶统马家沟组

下奥陶统马家沟组在鄂尔多斯盆地西南部中央古隆起和北部伊盟隆起缺失严重, 主要发育在鄂尔多斯

表1 鄂尔多斯盆地潜在氦源岩参数统计

Table 1 Parameter statistics of potential helium source rocks in the Ordos Basin

潜在层系	年龄/Ma	岩性	分布区	U 平均丰度*/10 ⁻⁶	Th 平均丰度*/10 ⁻⁶	生氦强度/ [10 ⁻⁶ cm ³ /(a·g)]
太古宇- 下元古界	230 ~ 250	黑云母片岩、 花岗片麻岩	北东-南西 向条带	(3.97 ~ 0.26)/1.29 (12)	(18.75 ~ 1.42)/8.19 (12)	0.391
中元古界 长城系	153.5 ~ 178.5	板岩	盆地北部、 西南部	(4.33 ~ 0.39)/1.29 (29)	(8.33 ~ 0.75)/2.49 (29)	0.227
下古生界下奥陶 统马家沟组	48.7	泥质白云岩	古陆周边	(8.98 ~ 1.53)/4.30 (8)	(25.61 ~ 2.28)/10.60 (8)	0.820
上古生界石炭系、 二叠系	28.5±5	泥岩	全盆地	(16.12 ~ 3.22)/9.68 (26)	(54.28 ~ 1.57)/22.68 (26)	1.820
		煤	全盆地	(40.49 ~ 2.87)/14.99 (15)	(88.12 ~ 2.00)/35.19 (15)	2.820
		铝土岩	盆地西南部	(11.85 ~ 1.35)/7.14 (11)	(59.39 ~ 6.33)/38.57 (11)	1.970

注: *数据格式为最小值 ~ 最大值/平均值(样本数)。

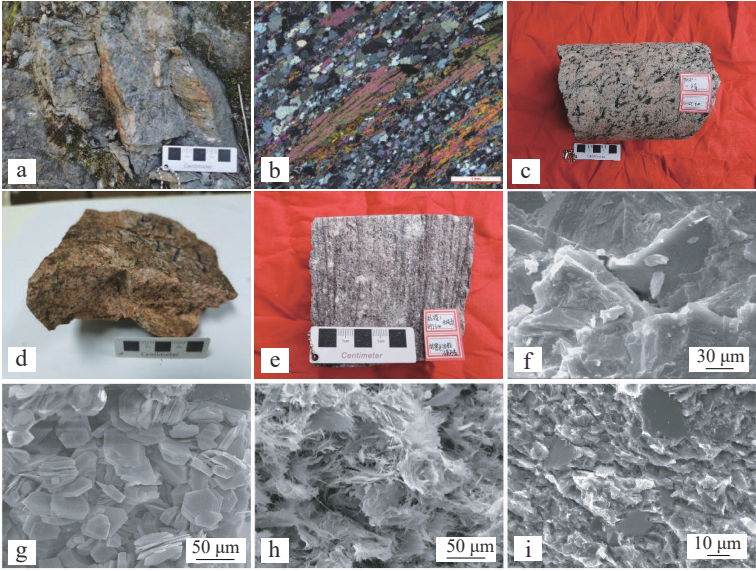


图3 鄂尔多斯盆地潜在氦源岩特征照片

Fig. 3 Micrographs and pictures showing potential helium source rocks in the Ordos Basin

a. 深色二云母石英片岩,北秦岭,太古宇,野外露头;b. 深色二云母石英片岩(+),北秦岭,太古宇,薄片;c. 花岗片麻岩,杭探2井,太古宇,埋深2 468.60 m,岩心;d. 紫苏花岗片麻岩,乌拉山-大青山,古元古界,野外露头;e. 石英砂岩,杭探1井,长城系,埋深3 973.00 m,岩心;f. 泥质白云岩,米109井,马五段,埋深2 410.82 m,SEM;g. 砂岩中的高岭石,米109井,本溪组,埋深2 396.88 m,SEM;h. 砂岩中的伊利石,米109井,太原组,埋深2 348.60 m,SEM;i. 铝土岩中的水铝石团块斑状分布,陇81井,太原组,埋深3 910.19 m,SEM

盆地中东部,根据沉积演化特征,马家沟组可以分为6个段:马一段、马三段和马五段主要为蒸发台地沉积,岩性主要为膏盐岩和白云岩;马二段、马四段和马六段主要为碳酸盐台地沉积,岩性主要为白云岩和石灰岩^[44]。以桃61井为例,马五段的白云岩夹有薄层泥质白云岩,厚度约为45 m,Th和U元素含量相对较高,具有一定的生氦潜力。

下古生界下奥陶统马五段泥质白云岩为潜在氦源岩(图3f)^[41],平均年龄为487 Ma,U和Th元素平均丰度分别为4.30×10⁻⁶和10.60×10⁻⁶,生氦强度为0.820×10⁻⁶ cm³/(a·g)。

3.3 石炭系-二叠系

上古生界石炭系-二叠系泥岩、炭质泥岩和煤是烃源岩,砂岩和铝土岩是天然气储层(图3g—i)。泥岩和煤在全盆地广泛分布,平均年龄为292 Ma,其中二叠系太原组泥岩U和Th元素平均丰度分别为9.68×10⁻⁶和22.68×10⁻⁶,泥岩厚度为30 m左右,生氦强度为1.820×10⁻⁶ cm³/(a·g);二叠系太原组煤U和Th元素平均丰度分别为14.99×10⁻⁶和35.19×10⁻⁶,最大厚度为14 m,生氦强度为2.820×10⁻⁶ cm³/(a·g)。由于相较于基底氦源岩,上古生界石炭系-二叠系泥岩和煤的厚度不大,制

岩,下古生界泥质含量较高的碳酸盐岩以及上古生界的泥岩。综合分析,生氢强度由强到弱依次为上古生界煤、上古生界铝土岩、上古生界泥岩、太古宇-古元古界花岗岩及变质岩、中-新元古界沉积变质岩和下古生界碳酸盐岩。但考虑到岩石年龄、氢源岩体积规模等氢源岩生氢潜力评价参数,对氢气生成的贡献仍需更深入的讨论。如太古宇-古元古界的花岗岩-变质岩系的厚度大,而沉积层系的厚度较小,例如石炭系-二叠系的煤系烃源岩厚度只有30~50 m,尽管其煤系烃源岩生氢强度更大,但单位时间生成的氢气总量远小于太古宇-古元古界的花岗岩-变质岩系。

受U和Th元素在氢源岩中的富集程度影响,鄂尔多斯盆地不同区域的生氢强度存在差异。整体上看,全盆地生氢强度在 $5.66 \times 10^{-6} \sim 9.28 \times 10^{-6} \text{ cm}^3/(\text{a} \cdot \text{g})$,盆地南部的生氢强度较高,具有较高的生氢潜力(图4)。其中盆地西南部庆阳气田的氢源岩生氢强度最大,为 $9.28 \times 10^{-6} \text{ cm}^3/(\text{a} \cdot \text{g})$,而盆地北部东胜气田的氢源岩生氢强度最小,为 $5.66 \times 10^{-6} \text{ cm}^3/(\text{a} \cdot \text{g})$ 。相较于鄂尔多斯盆地其他地区,各类潜在氢源岩在鄂尔多斯盆地西南部均有发育,具有“多源供氢”特征,氢源岩的体积规模优势较为明显,因此鄂尔多斯盆地西南部生氢潜力最大。

5 结论

1) 鄂尔多斯盆地潜在氢源岩为地壳型,包括基底岩系和沉积岩系,主要是太古宇-古元古界富含U-Th元素的花岗岩和花岗片麻岩等变质岩、中元古界长城系沉积变质岩、下古生界泥质含量较高的碳酸盐岩及上古生界的泥岩、煤和铝土岩。

2) 太古宇和古元古界基底岩系生氢强度为 $0.391 \times 10^{-6} \text{ cm}^3/(\text{a} \cdot \text{g})$,长城系板岩、下奥陶统马家沟组泥质白云岩及石炭系-二叠系泥岩、煤和铝土岩生氢强度分别为 $0.227 \times 10^{-6} \text{ cm}^3/(\text{a} \cdot \text{g})$, $0.820 \times 10^{-6} \text{ cm}^3/(\text{a} \cdot \text{g})$, $1.820 \times 10^{-6} \text{ cm}^3/(\text{a} \cdot \text{g})$, $2.820 \times 10^{-6} \text{ cm}^3/(\text{a} \cdot \text{g})$ 和 $1.970 \times 10^{-6} \text{ cm}^3/(\text{a} \cdot \text{g})$ 。

3) 各类潜在氢源岩在鄂尔多斯盆地西南部均有发育,具有“多源供氢”特征,生氢强度大,是潜在的富氢天然气分布区。

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